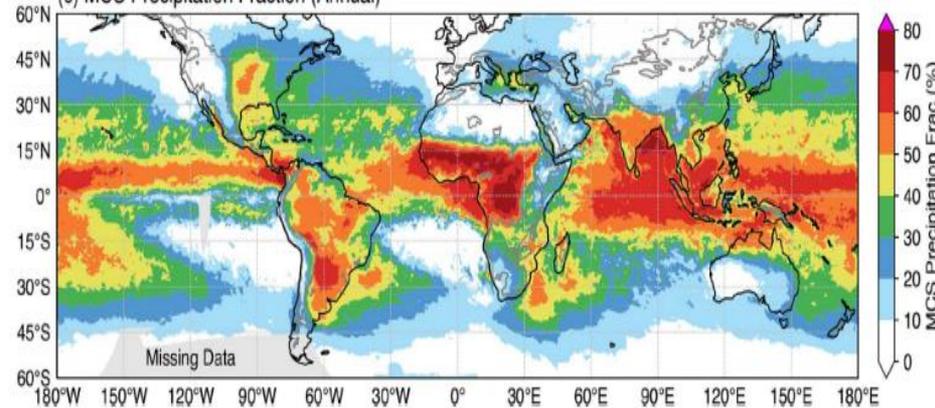
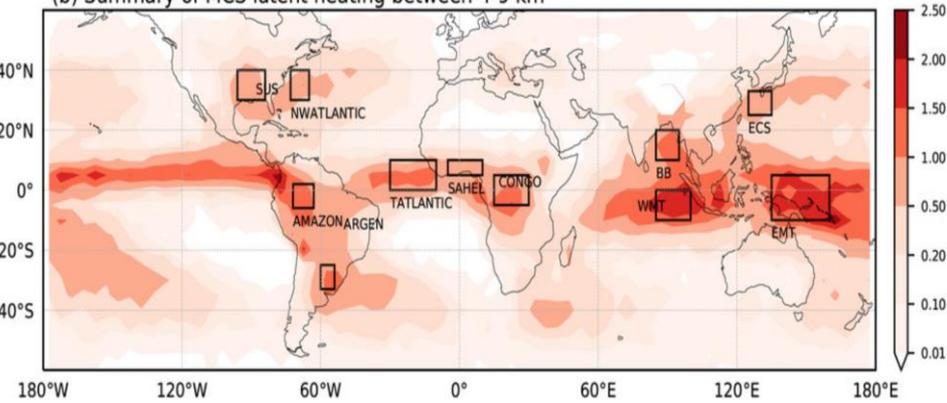


# Towards a stochastic parameterization of unresolved MCS impacts on large scales

(c) MCS Precipitation Fraction (Annual)



(b) Summary of MCS latent heating between 4-9 km



Robert Plant<sup>1</sup>, Hannah Christensen<sup>2</sup>, Mark Muetzfeldt<sup>1</sup>, Tim Woollings<sup>2</sup>, Zhixiao Zhang<sup>2</sup>

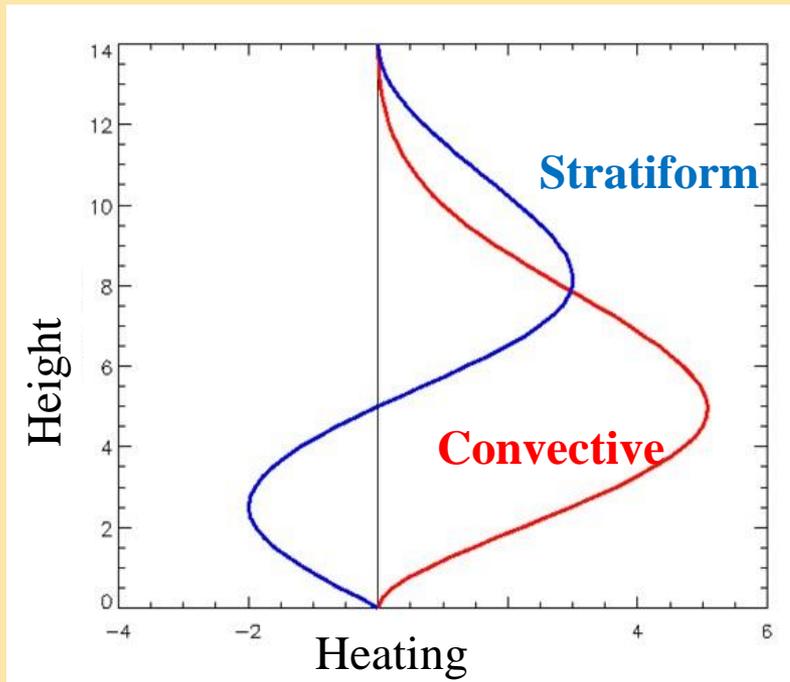
1. Department of Meteorology, University of Reading, UK

2. Atmospheric, Oceanic and Planetary Physics, University of Oxford, UK

AOGS2023, Singapore  
30 July - 4 August 2023

# Introduction

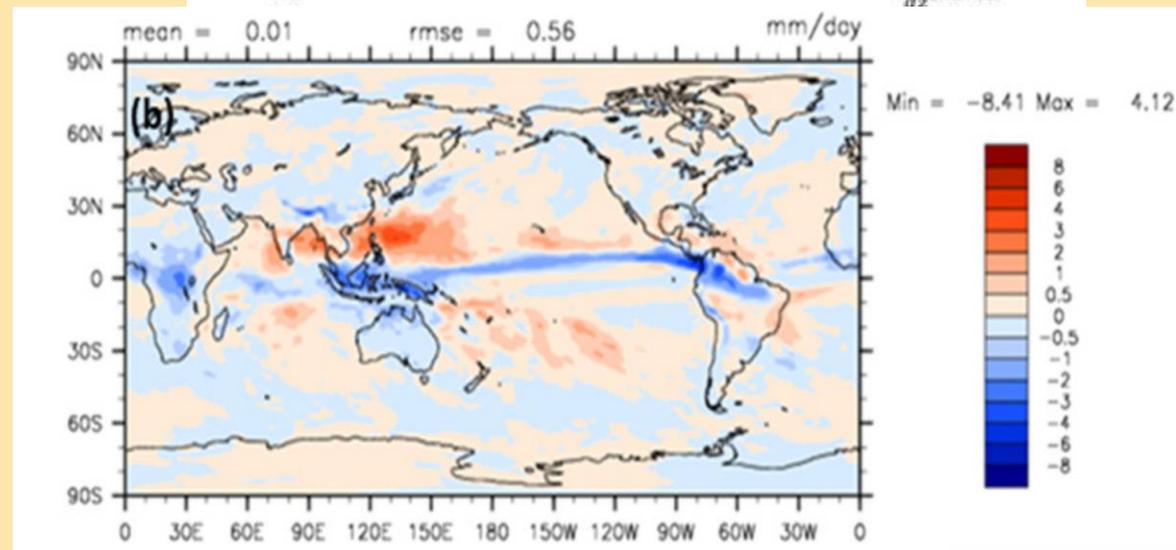
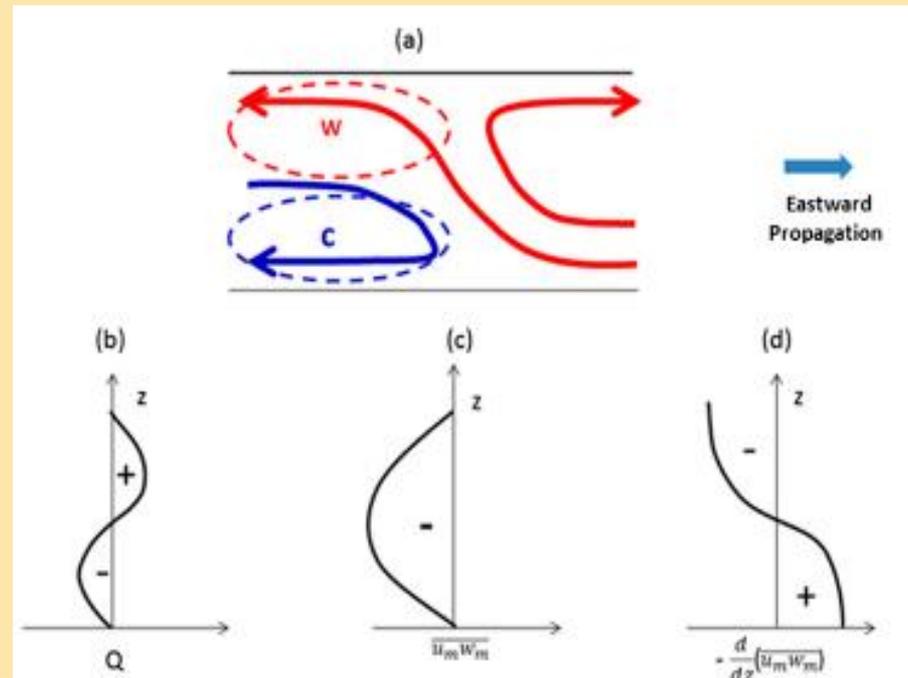
- Mesoscale Convective Systems responsible for >50% of rainfall in large areas of tropics
- MCSs alter convective heating profile. Can couple to dynamics and reshape large-scale circulations



- Heating aloft, but cooling in lower troposphere, where melting and evaporation occur
- Global models parameterize convection
- And struggle to represent the stratiform part

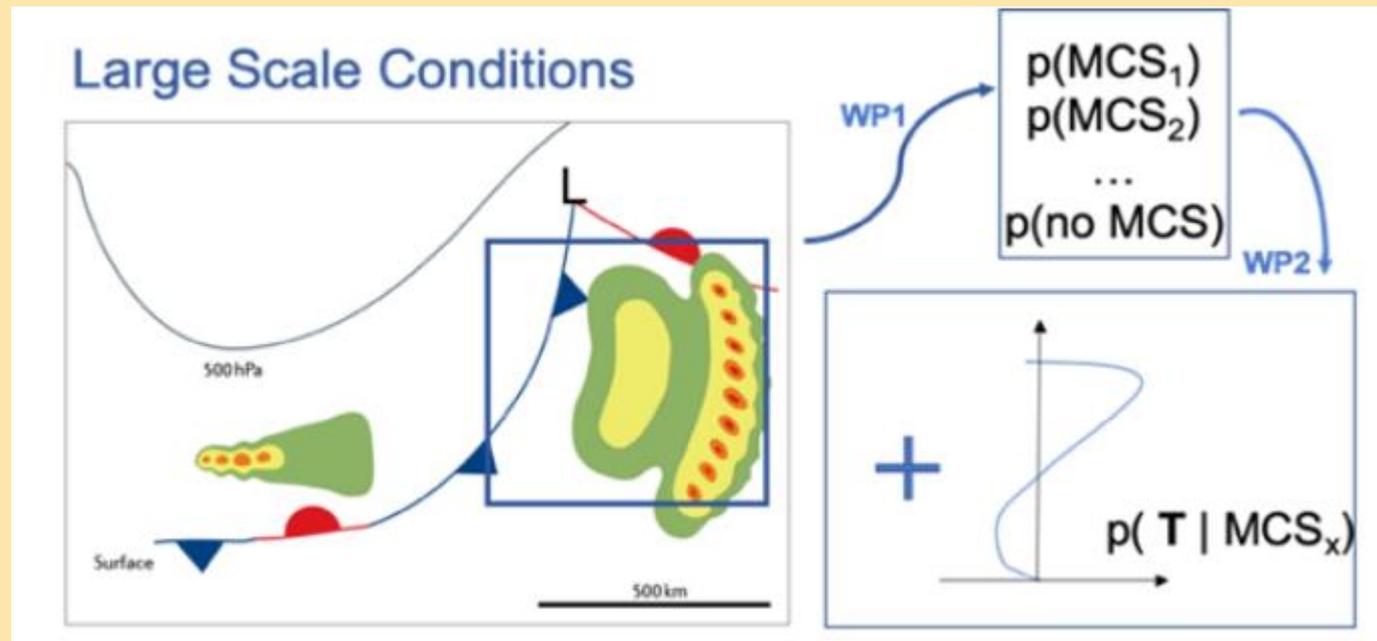
# Parameterizing MCS effects?

- Multiscale Coherent Structure Parameterization (Moncrieff et al 2017) makes heating profile more top-heavy and provides additional momentum transports
- Improved precipitation across the ITCZ and Maritime Continent in CAM and E3SM

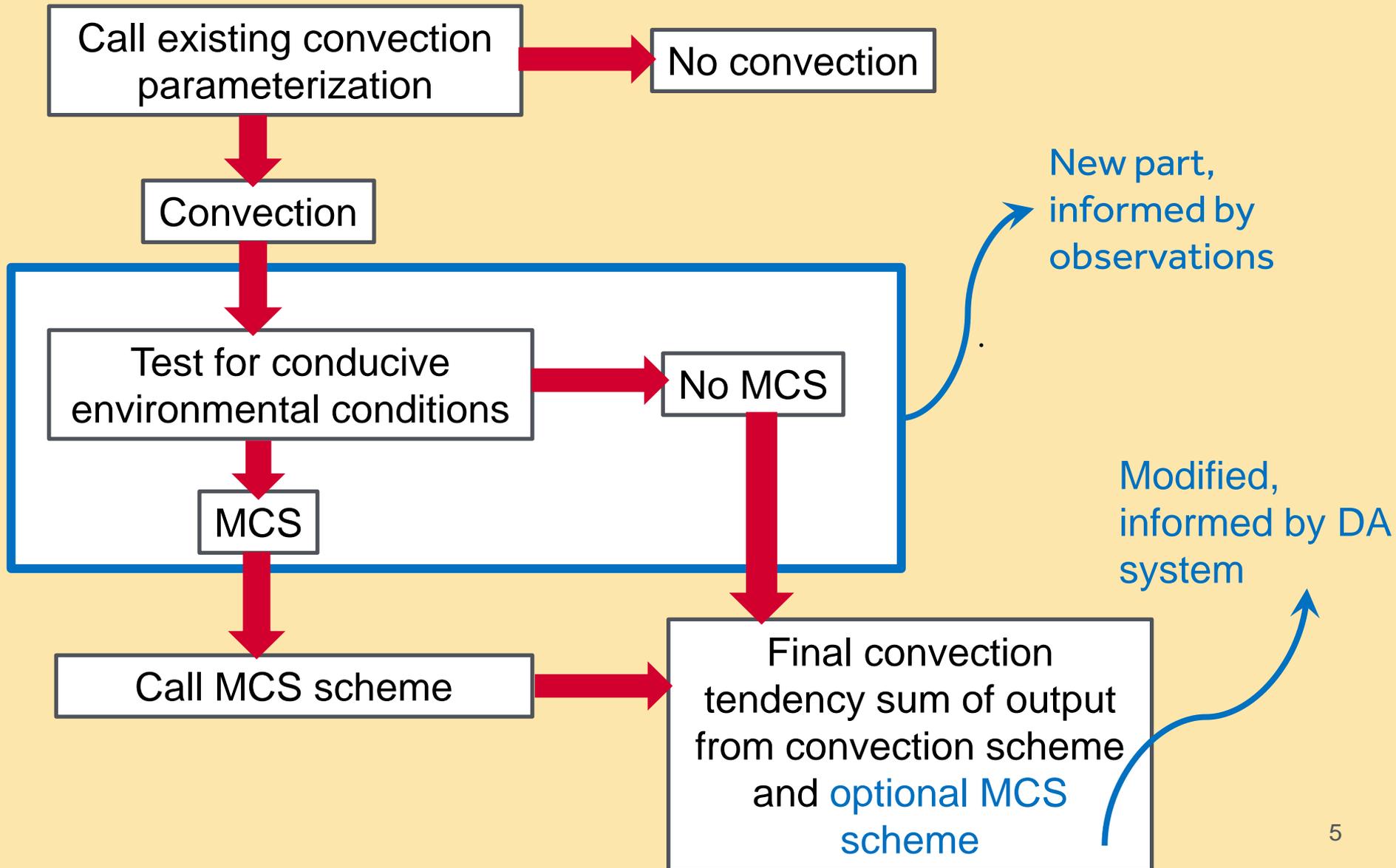


# MCS:PRIME

- We target global ensemble NWP
- Develop the ideas by:
  - including dependence of occurrence on environmental state
  - reformulating missing tendencies based on analysis of DA increments
- Both parts with a stochastic formulation



# MCS scheme envisaged

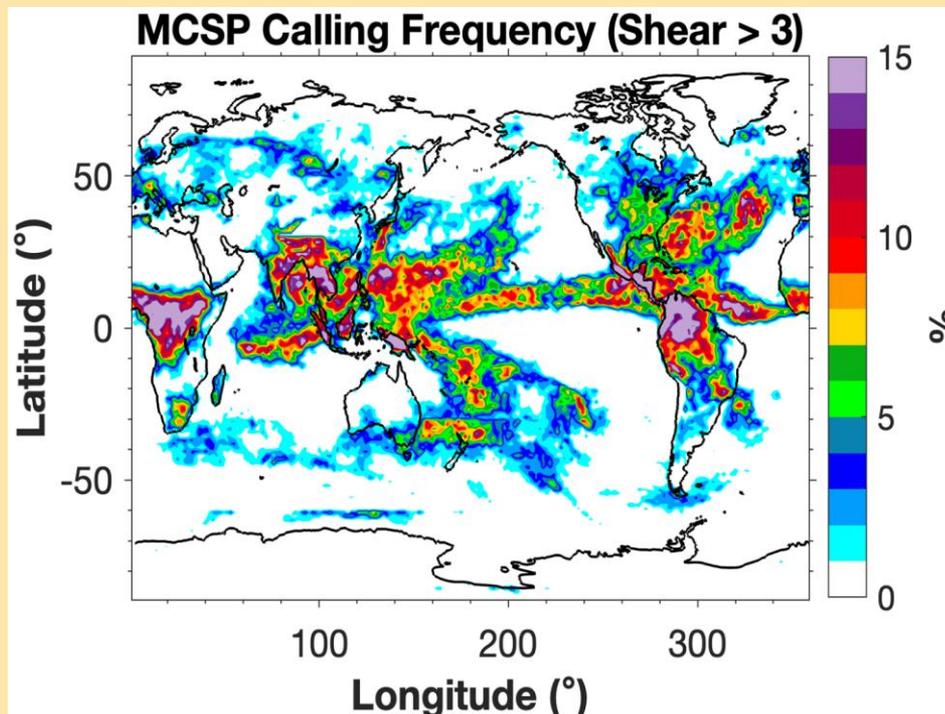


# CoMorph-A Convection Scheme

- New mass flux scheme for the Unified Model
- Competitive but not yet operational
- Initiated at any level, proportional to buoyant instability
- No distinction between shallow, deep and mid-level
- Entrainment rate depends on the turbulent mixing-length in the parcel's source layer, and on precipitation in the previous timestep (i.e., simple memory effect)
- Separate consideration of cloud-mean and cloud-core properties in detrainment calculations
- Implicit numerics prevent artificial on/off noise
- Developed through large Met Office / university partnership, ParaCon, but especially by Mike Whittall

# A minimal implementation

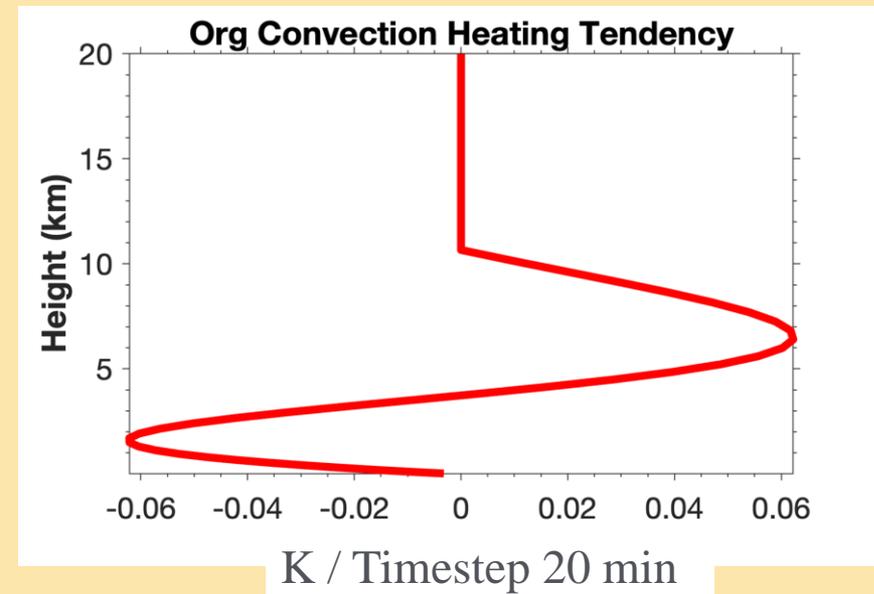
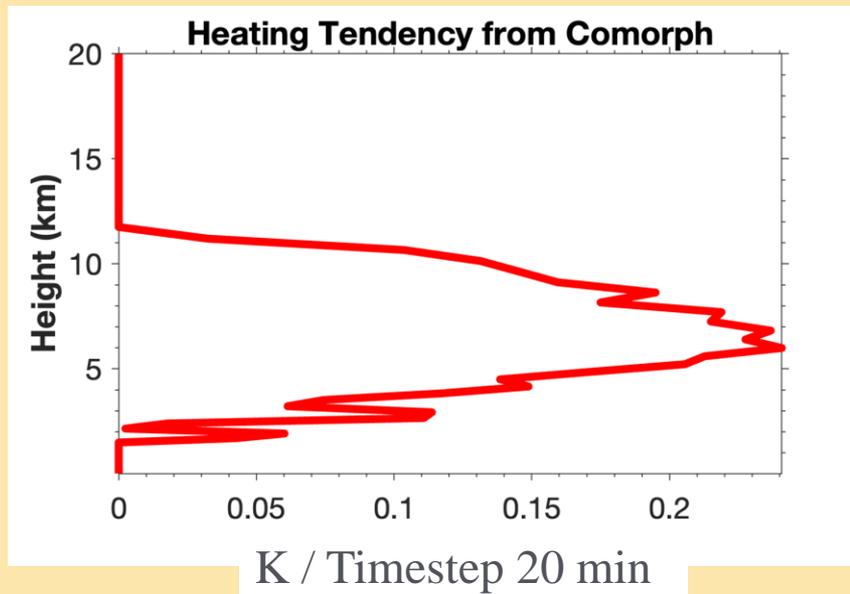
- MCS requires deep convection
  - Cloud top temperature  $< 0^{\circ}\text{C}$
  - Cloud base pressure  $> 600$  mb
  - Cloud base and top pressure difference  $> 300$  mb
- And requires some wind shear
  - Speed difference (600mb – lowest level)  $> 3$   $\text{ms}^{-1}$



Fraction of  
timesteps where  
MCS scheme  
active

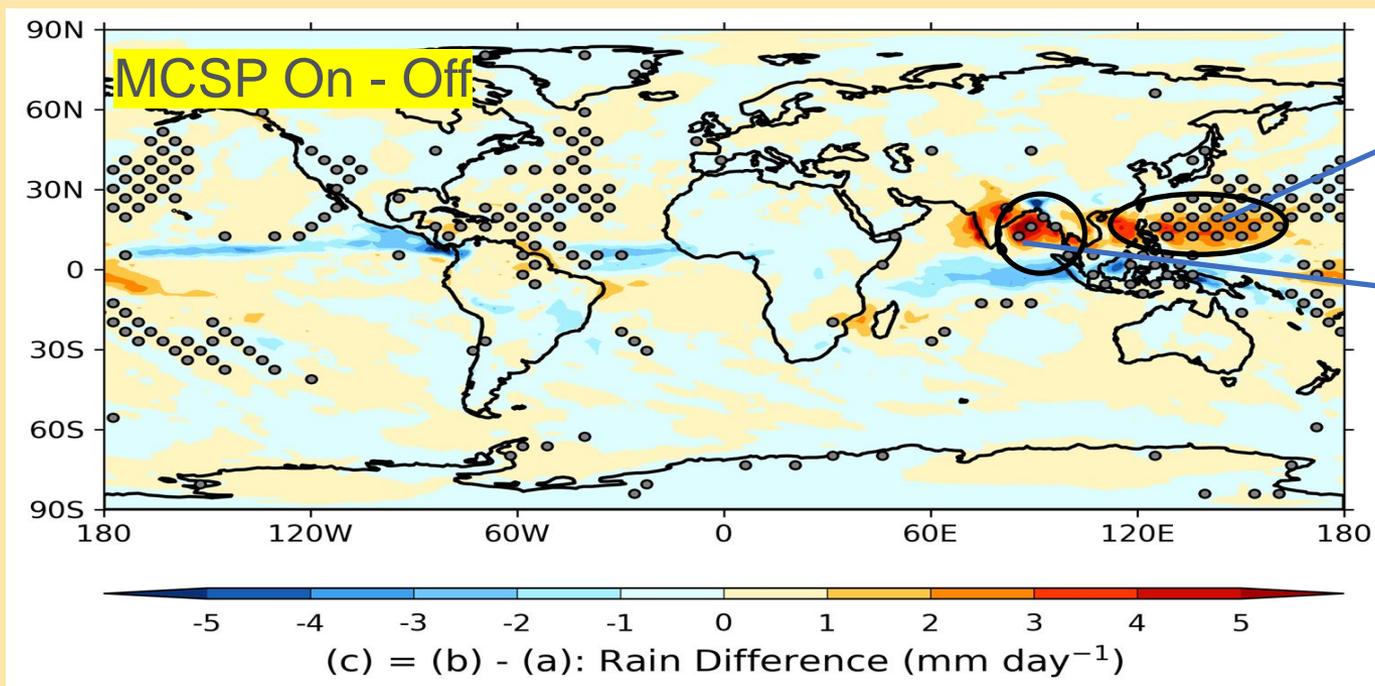
# A minimal implementation

- An example
- Peak value is  $\sim 25\%$  of CoMorph peak



# A 5 year climate test run

- Following AMIP-UM setup in CMIP6
- N96 (192x144 points, 85 vertical levels,  $\Delta t=20\text{min}$ )



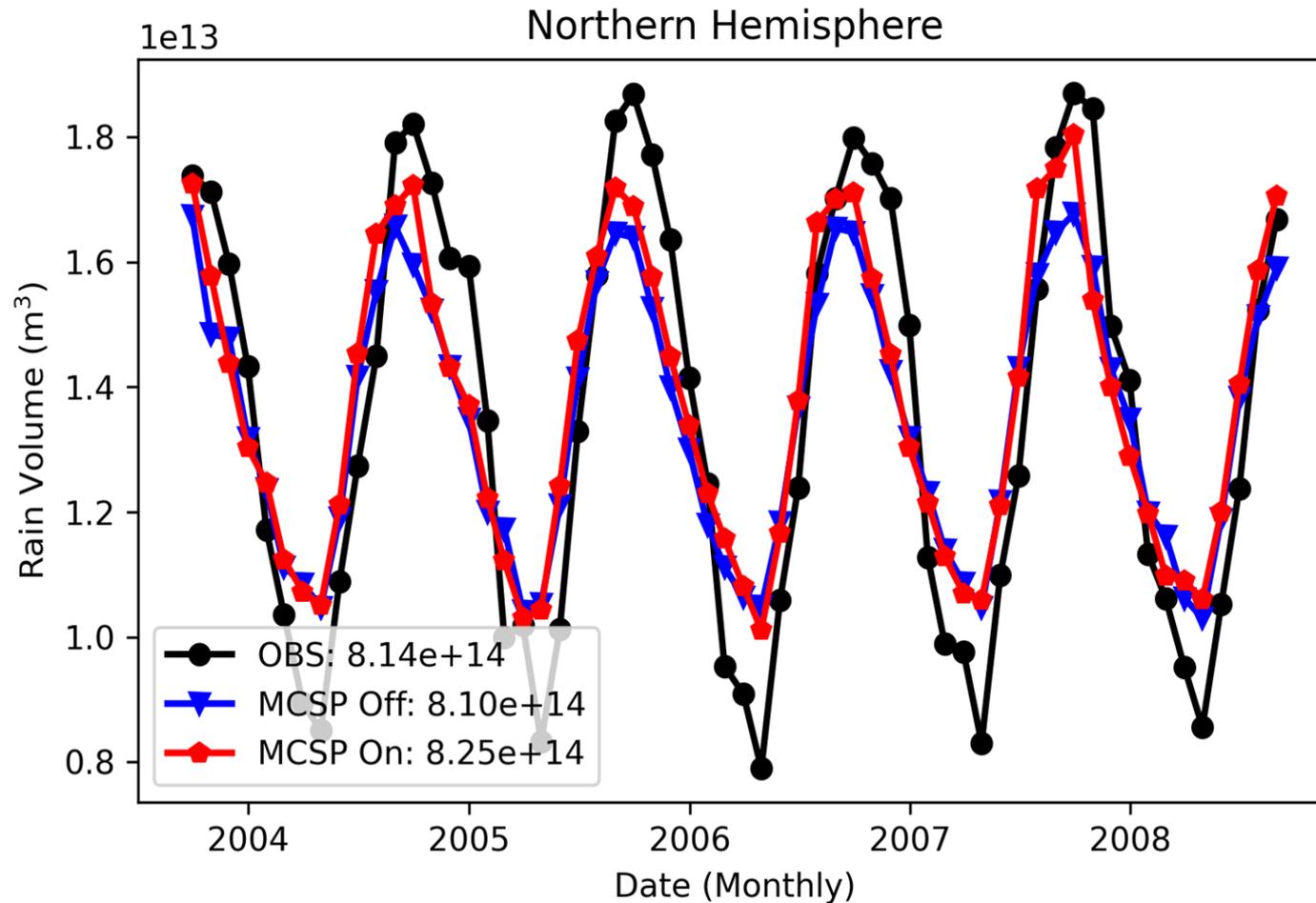
Increase of wet bias

Decrease of dry bias

Dots where this UM run significantly disagrees with CMIP6 (outside 5-95 percentile range)

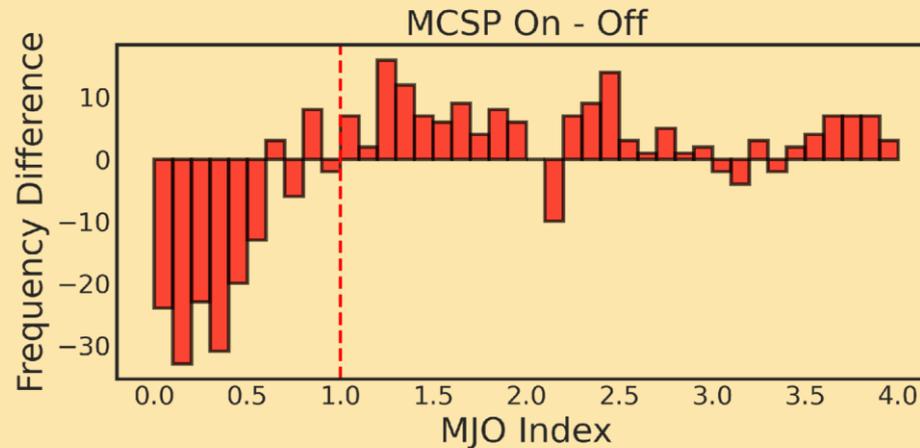
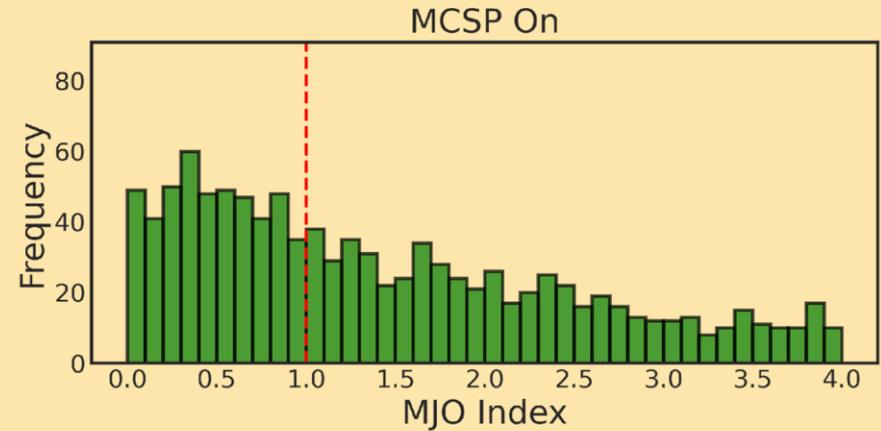
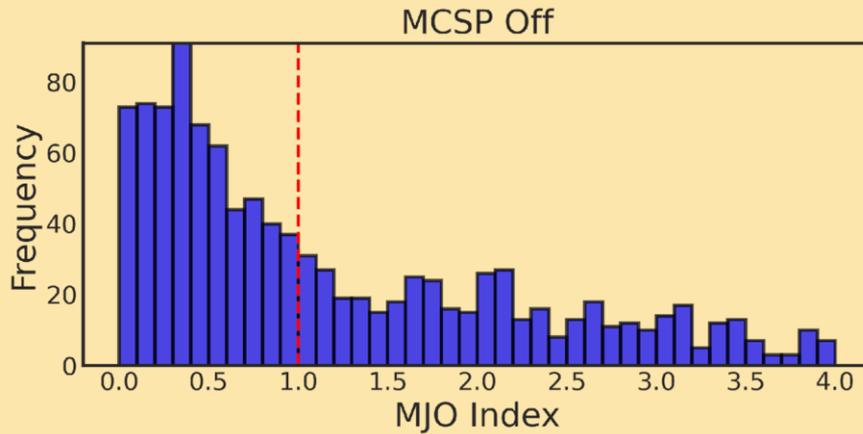
- Broadly similar to CAM and E3SM
- Emphasizes changes with CoMorph compared to old UM scheme

# Precipitation seasonal cycle



Stronger N hemisphere cycle due to **stronger Asian monsoon**

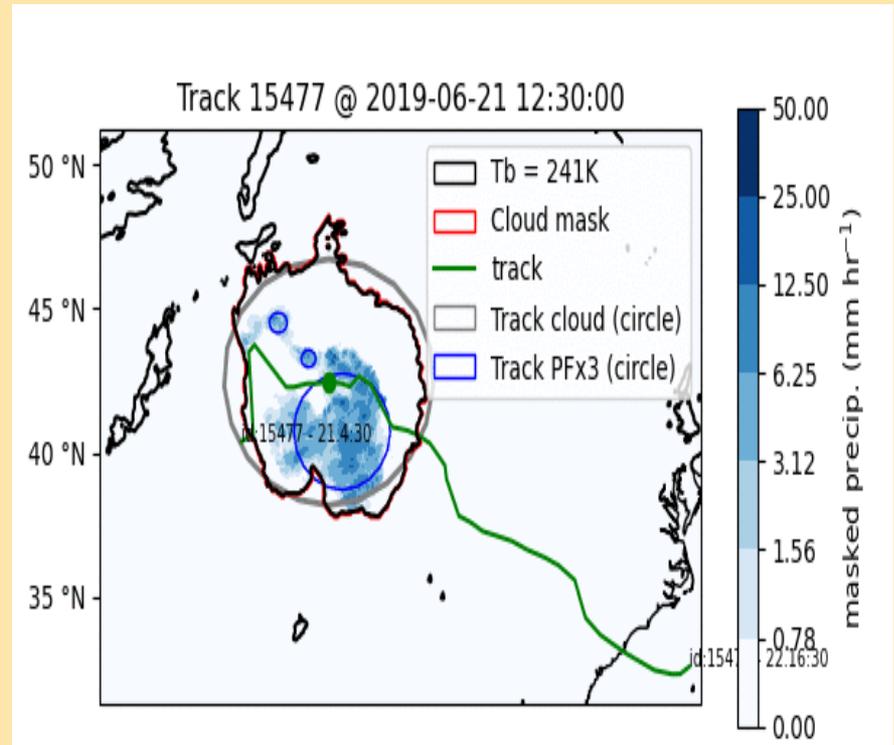
# Enhanced MJO



Relative stabilization at low levels discourages stationarity / encourages eastward progression of convection

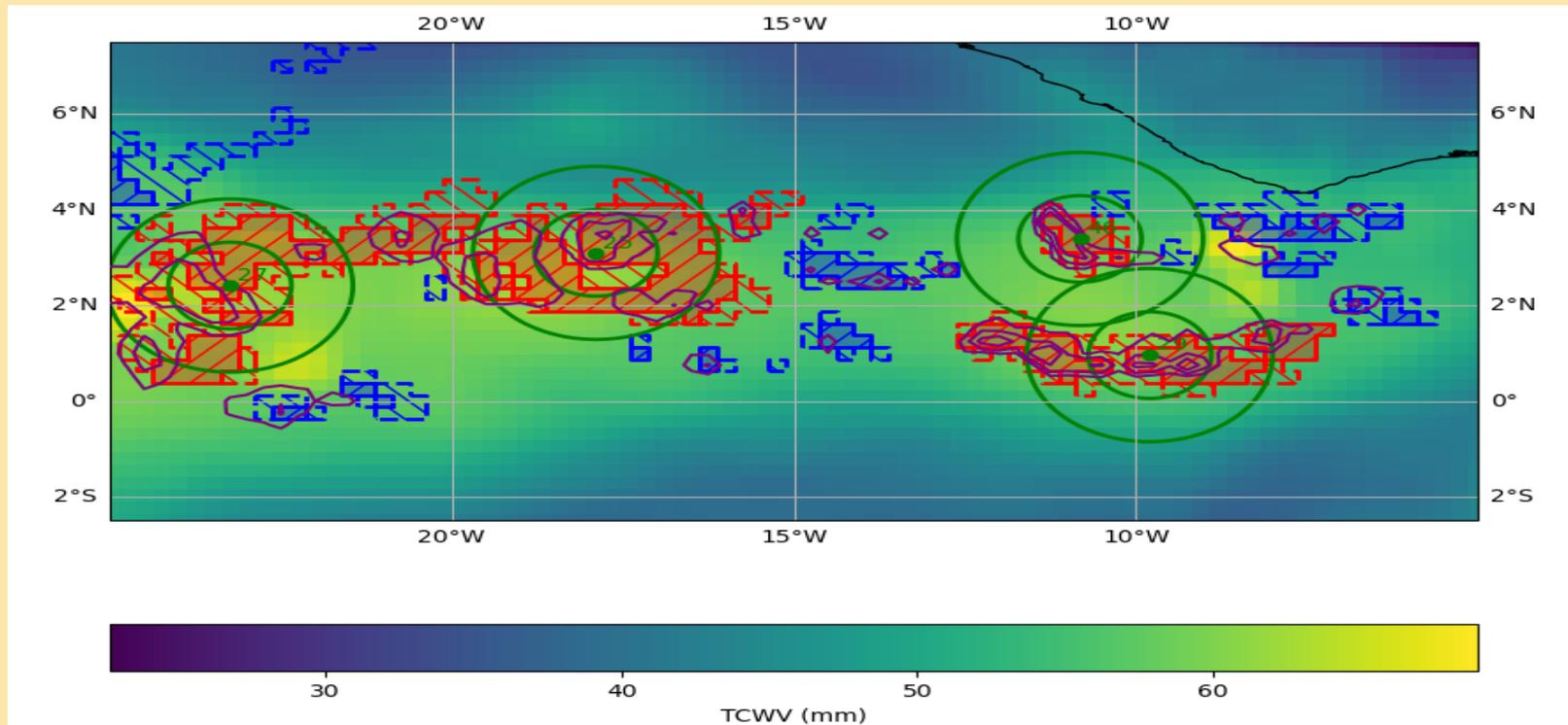
# Observational tracking data

- Feng et al. (2021) data from 2000-2020
- Based on NASA Global Merged IR V1 infrared brightness temperature,  $T_b < 225$  K for cloud core,  $T_b < 241$  K for cloud shield, as well as IMERG precipitation
- MCS area  $> 4 \times 10^4 \text{km}^2$ , duration  $> 4$  h, as well as other lifetime-dependent thresholds



Match to ERA5 conditions for CAPE, Total Column Water Vapour (TCWV), vertically integrated Moisture Flux Convergence (MFC), and various measures of shear

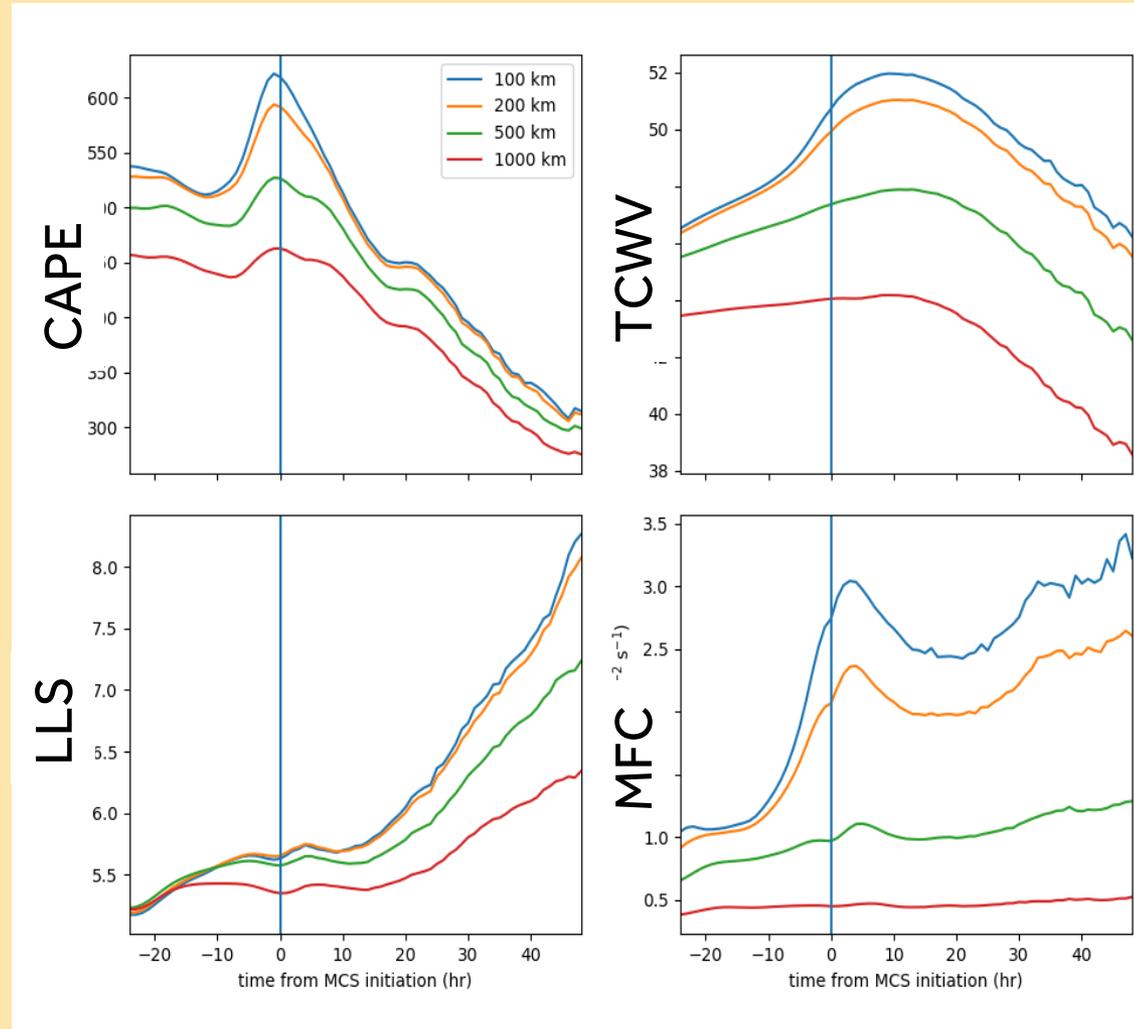
# Analysis regions overlaid on TCWV



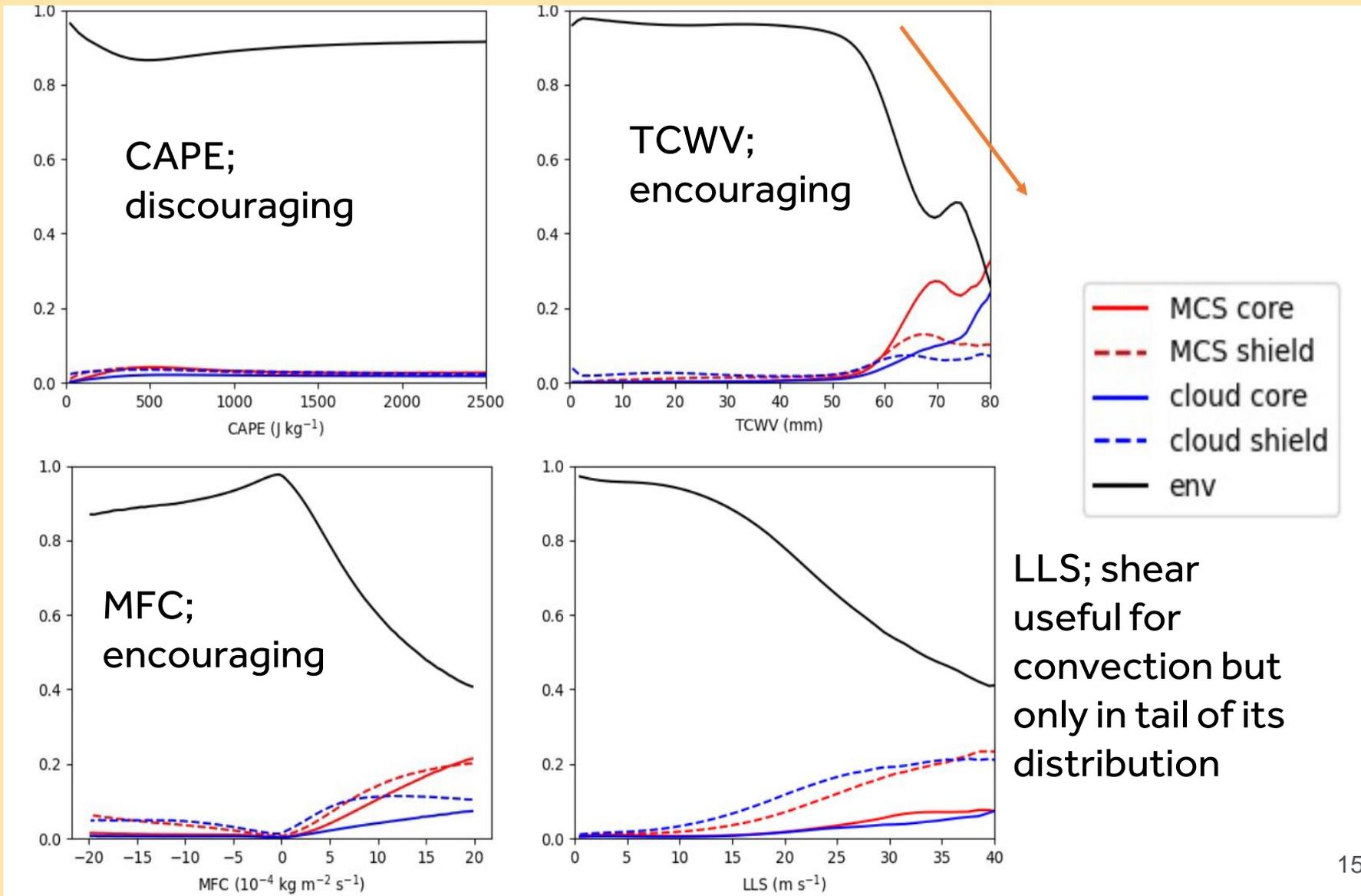
- Shield < 241K (dashed)
- Core < 225K (solid)
- Red=MCS
- Blue=convective, non-MCS
- Green circles: 100, 200km

# Evolution of MCS environment

- Before MCS initiation, use initiation centroid
- After MCS initiation, use track centroid
- Increases in CAPE, TCWV and MFC 5-10 hr beforehand. Smallest spatial scale has the largest change.
- Low-level shear shows a weak increase, but stronger increase over the next 15h
- $P(\text{environment} \mid \text{MCS})$  is interesting but....

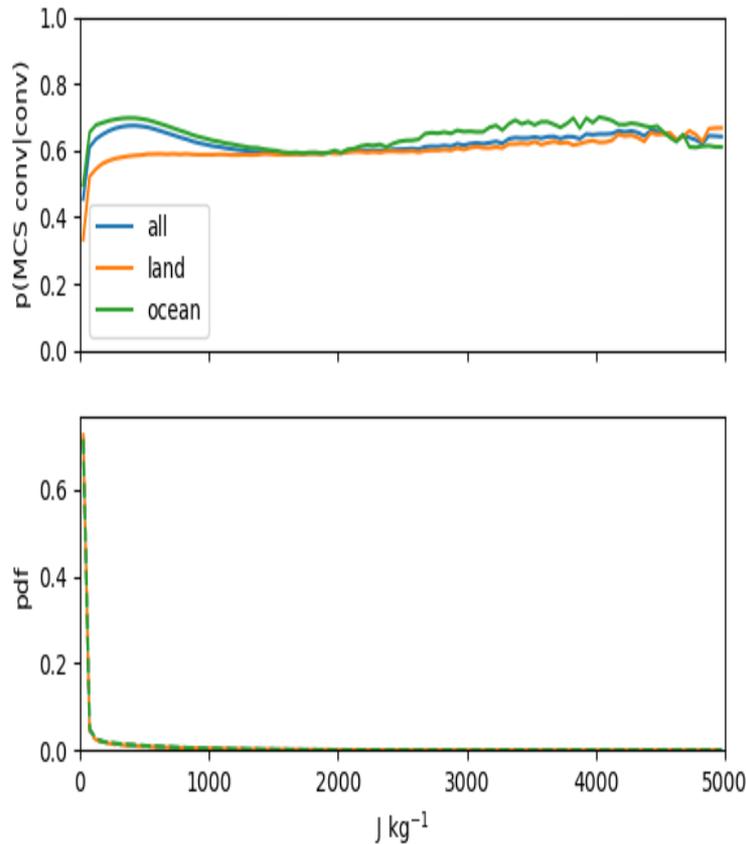


# Want to know P (MCS | variable)

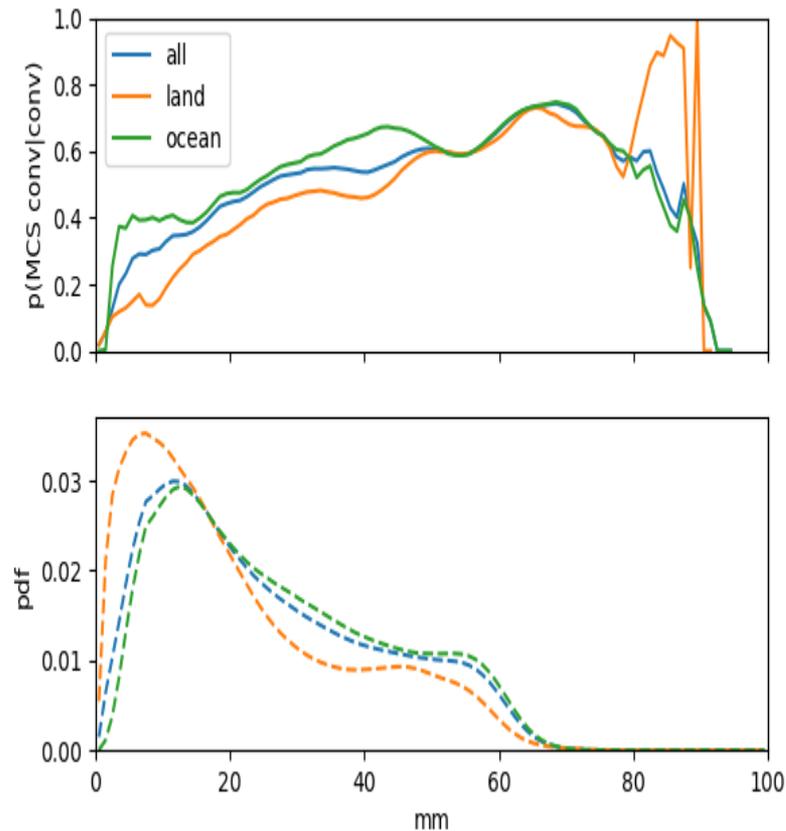


# Scheme needs P (MCS | conv, variable)

## CAPE



## TCWV



Condi  
tional  
probs

prob of  
variable

- Clear dependence on TCWV, indicating route for including environmental conditions in a stochastic call of the MCS parametrization

# Conclusions

- **A parameterization of MCS effects has substantial potential** to improve global NWP and climate models
  - Improved seasonal cycle in N hemisphere and the MJO
  - Reduced bias over Bay of Bengal
  - But amplifies a bias over the western Pacific
- Clear increases in different environmental variables 5-10 hours before MCS initiation
  - **There is precursor information in the environment**
- **Decision to activate the MCS is based on observed  $P(\text{MCS} | \text{convection, env state})$** 
  - Shear has little predictive power from this perspective
  - **TCWV has the clearest signal for distinguishing MCS**
- Next steps
  - Implement MCS predictors
  - Revise MCS tendencies with guidance from DA increments associated with MCS

# References

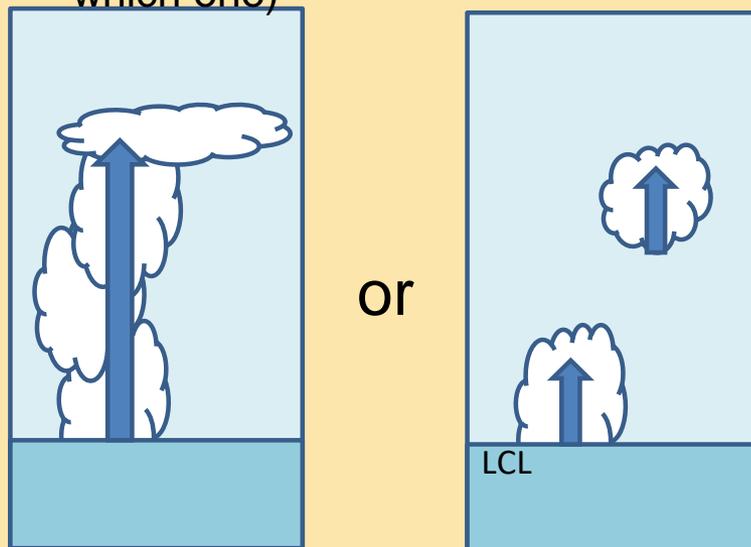
- Chen, C.-C., Richter, J. H., Liu, C., Moncrieff, M. W., Tang, Q., Lin, W., et al., 2021: Effects of organized convection parameterization on the MJO and precipitation in E3SMv1. Part I: Mesoscale heating. *J. Adv. Model. Earth Syst.*, **13**, e2020MS002401
- Daleu, C. L., Plant, R. S., Stirling, A. J. and Whittall, M. A. W, 2021. Evaluating the CoMorph parameterization using idealised simulations of the two-way coupling between convection and large-scale dynamics. To appear in: *Q. J. R. Meteorol. Soc.*, 2023.
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- Houze, R. A., 2004: Mesoscale convective systems. *Rev. Geophys.*, **42**, RG4003
- Liu, N., L. R. Leung, and Z. Feng, 2021: Global Mesoscale Convective System Latent Heating Characteristics from GPM Retrievals and an MCS Tracking Dataset. *J. Climate*, **34**, 8599–8613.
- Moncrieff, M. W., C. Liu, and P. Bogenschutz, 2017: Simulation, Modeling, and Dynamically Based Parameterization of Organized Tropical Convection for Global Climate Models. *J. Atmos. Sci.*, **74**, 1363–1380.

# Extras

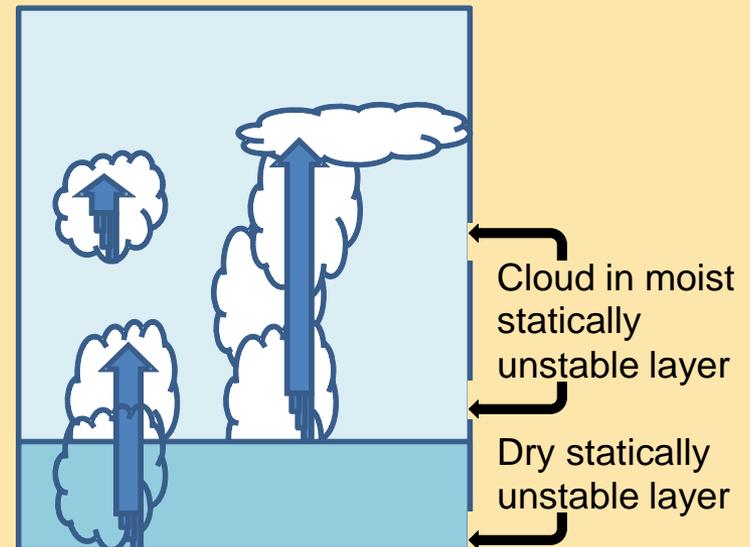
- CoMorph formulation
- Intermittency in CoMorph
- Rain rate plots
- Precipitation seasonal cycle in S hemisphere
- MJO related plots
- Conditional probabilities for MFC

## “Traditional” approach:

- Complex empirical trigger functions
- A-priori diagnosis of a unique “cloud-base” height
- Plume can only start from surface or other prescribed height
- Separate schemes for “deep”, “shallow” and “mid-level” convection (must pre-diagnosed which one)



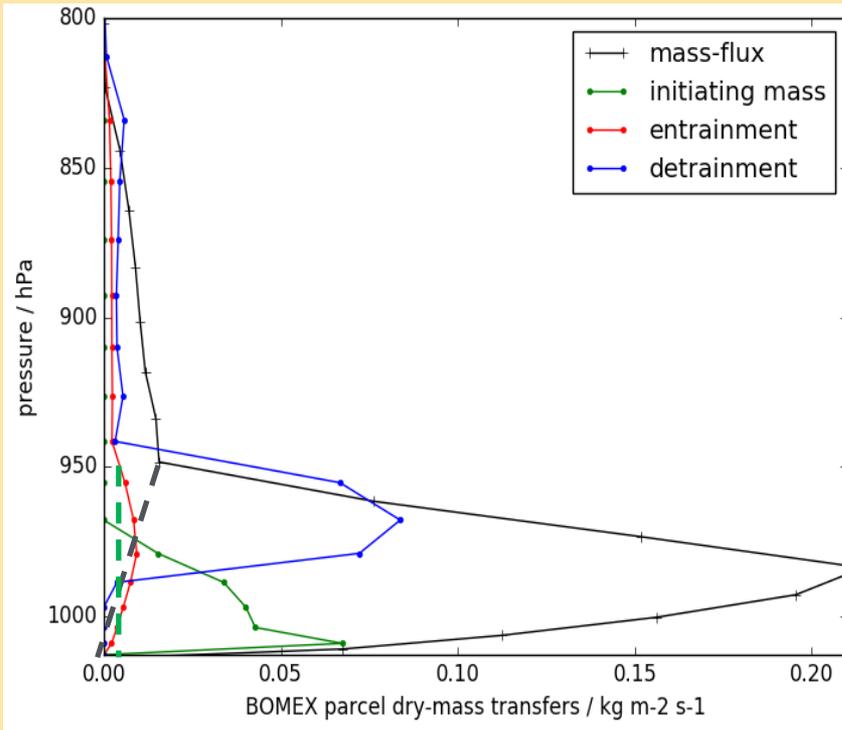
- Convecting parcels launch from any height where there is local vertical instability
- Plumes from different unstable layers integrated independently
- Same parcel ascent / descent code for all plumes
- Updraft radius depends on the turbulent mixing-length in the parcel’s source-layer.



# CoMorph Mass Flux Budget

## Diagnostic bulk vertical mass transport budget

$$\frac{\partial M}{\partial z} = g + e - d$$



## Initiating mass-source

depends on local instability:  
(dry-static instability in clear-air, moist-static instability in large-scale cloud)

$$g = \frac{1}{4} \rho \sqrt{-N^2}$$

## Entrainment rate

inversely proportional to radius. Radius is proportional to turbulence length-scale from the boundary-layer scheme, with enhancement by the previous precip

$$e = M \frac{0.2}{R}$$

$$R_{init} = \alpha h_{BL}$$

$$\alpha = \alpha_0 + \frac{pr/q}{pr_{max}/q_{ref}} (\alpha_{max} - \alpha_0)$$

## Detrainment rate

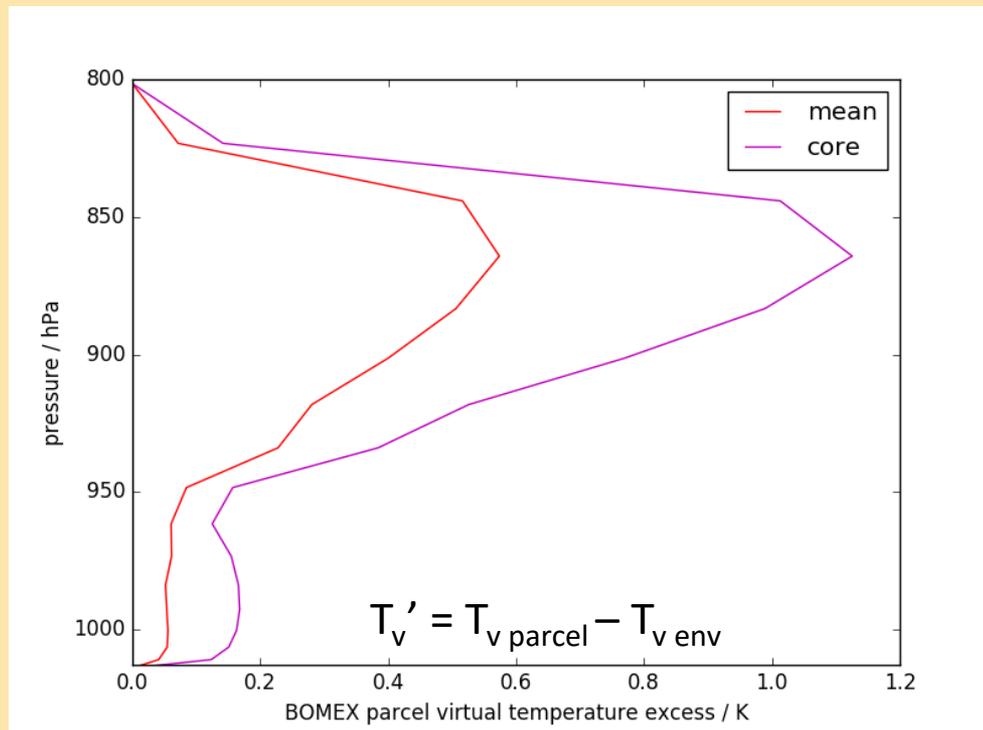
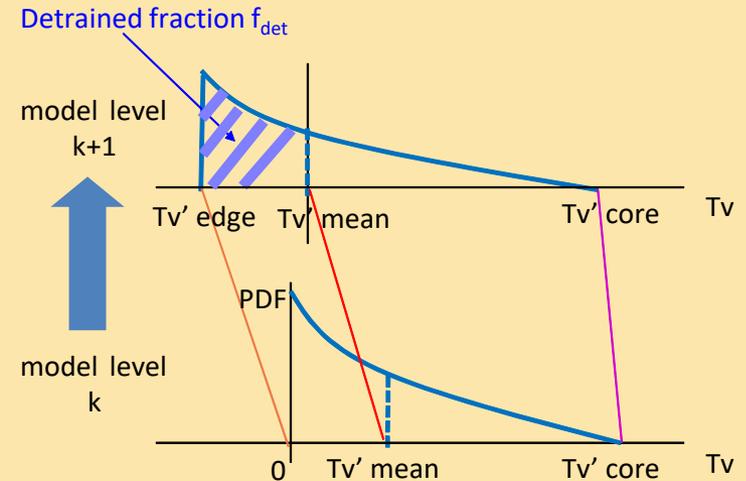
based on sorting an assumed-PDF of buoyancy within the bulk plume, with implicit numerical method

$$d = \frac{M}{dz} \int_{T_{vedge}}^{T_{venv}} PDF(T_v) dT_v$$

# CoMorph Detrainment

## Implicit Assumed-PDF-based detrainment

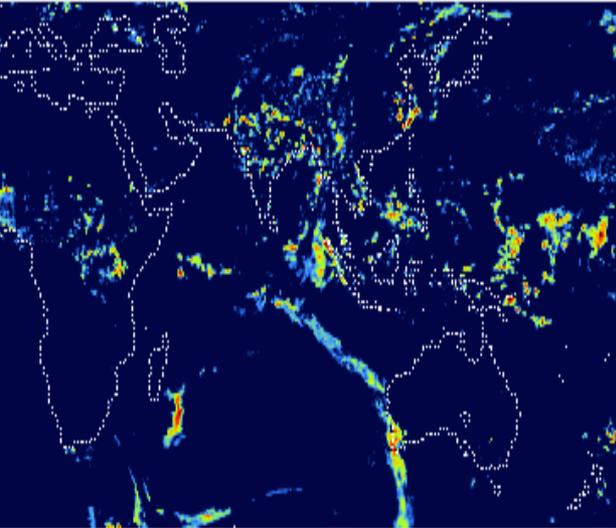
- Buoyancy,  $T$ ,  $q$ ,  $u$ ,  $v$ , etc assumed to have a power-law PDF *within the bulk plume*.
- Separate ascent calculations for in-plume mean properties, and a less dilute parcel “core”
- At each level, detrain the fraction of the PDF which becomes non-buoyant (due to entrainment, changes in environment  $T_v$ , etc)



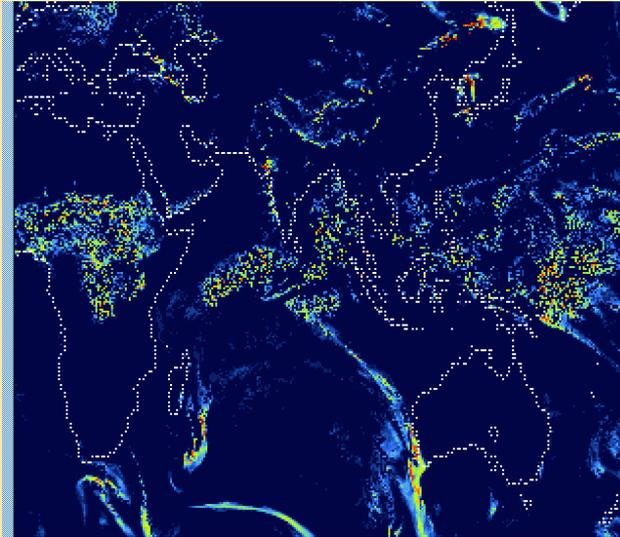
But how to define  $T_{v \text{ env}}$  when it is changing due to the convective increment?

For smooth behaviour. need implicit-in-time discretization, accounting for the convective heating.

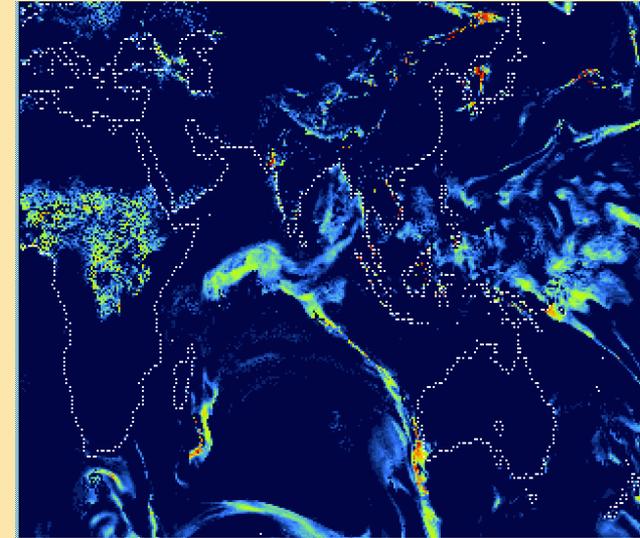
# A snapshot of rain rates



TRMM data



Old scheme,  
CAPE-based  
closure

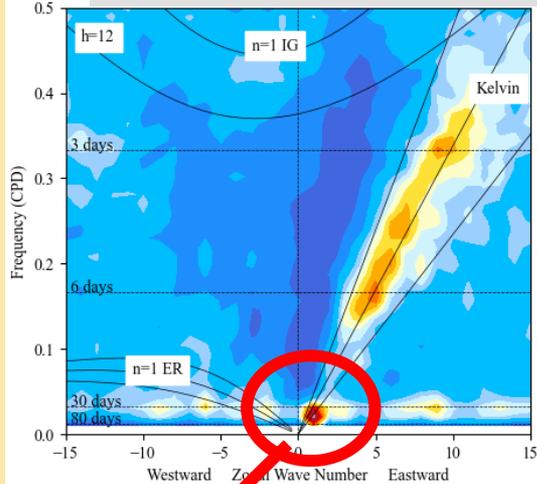


A CoMorph test

# Equatorial waves with CoMorph

- N96, atmosphere only.
- Tropical wavenumber-frequency spectrum for precipitation

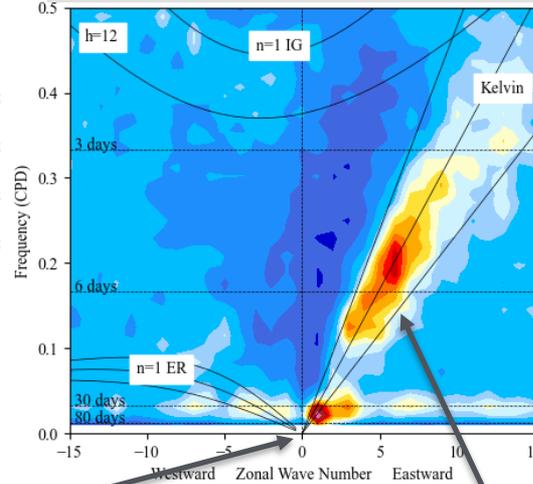
Ctrl (GAL9)



MJO

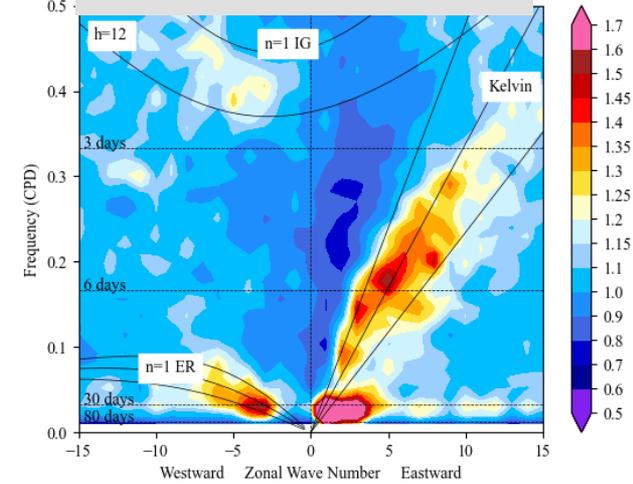
Slightly stronger  
MJO

CoMA9



Improved Kelvin wave  
strength

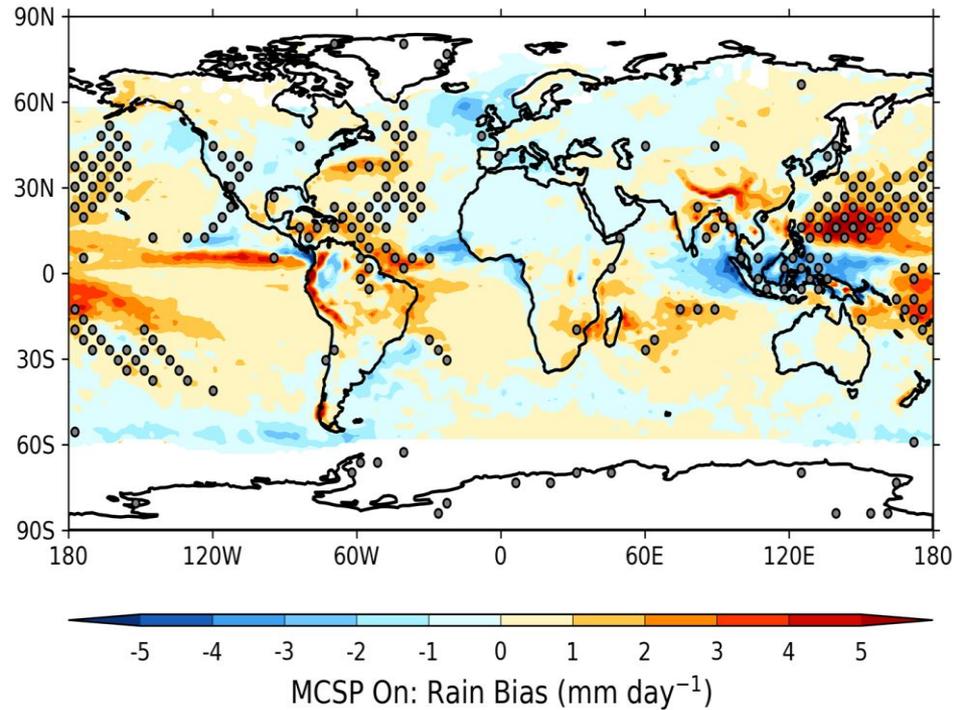
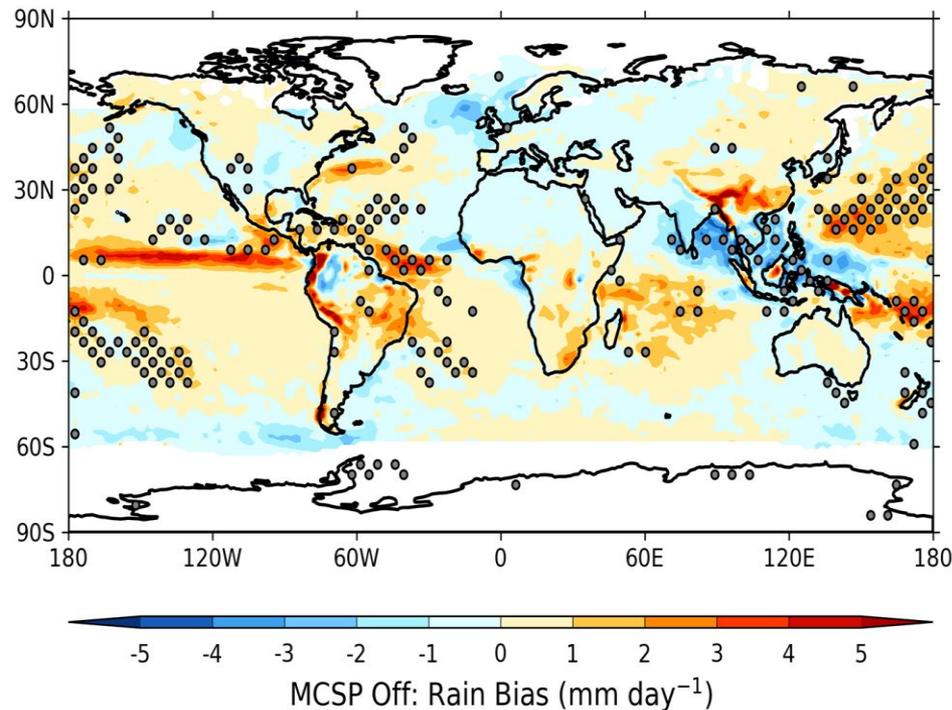
Observations



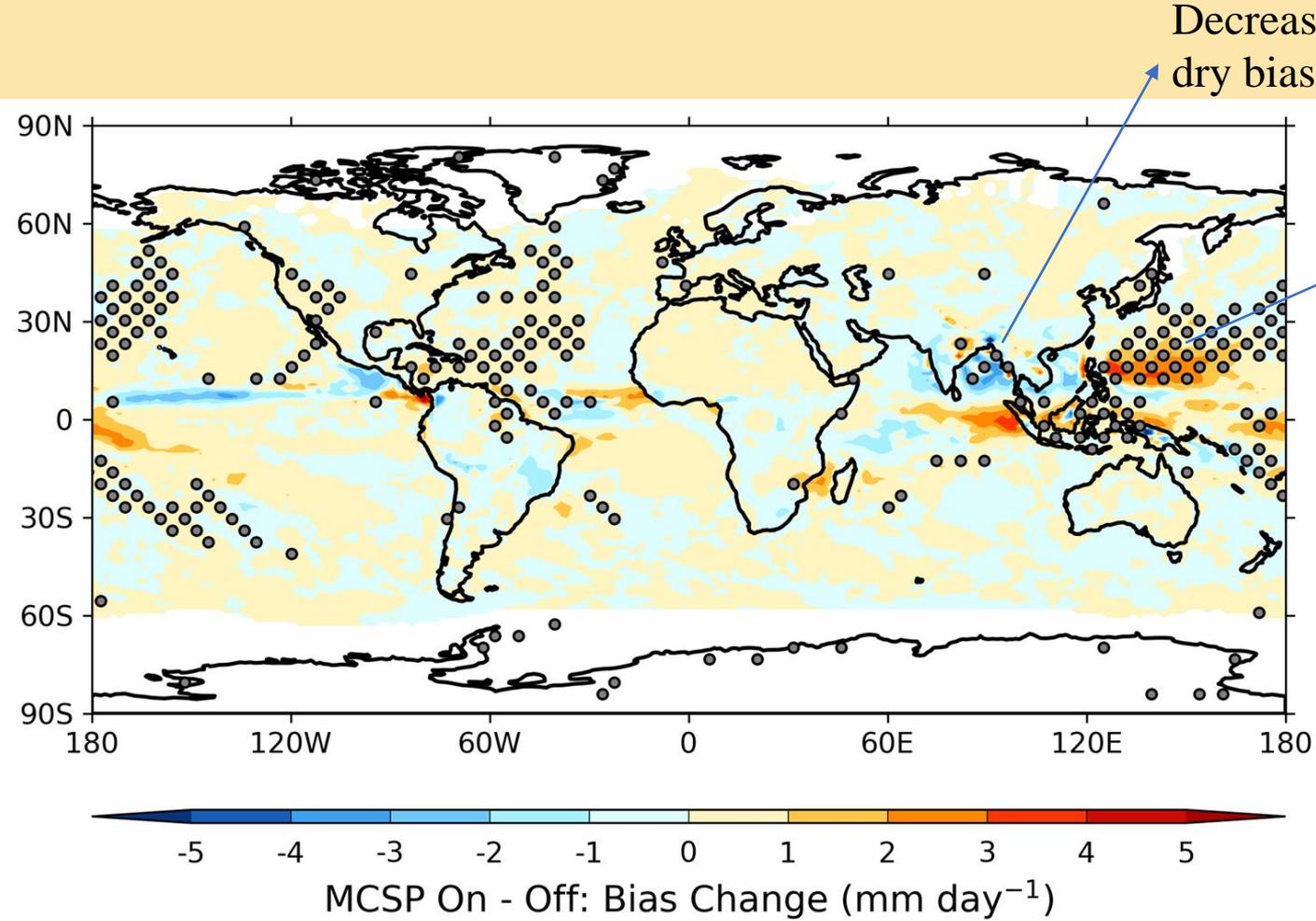
# Rain rate bias

MCSP Off- OBS  
correlation: 0.886

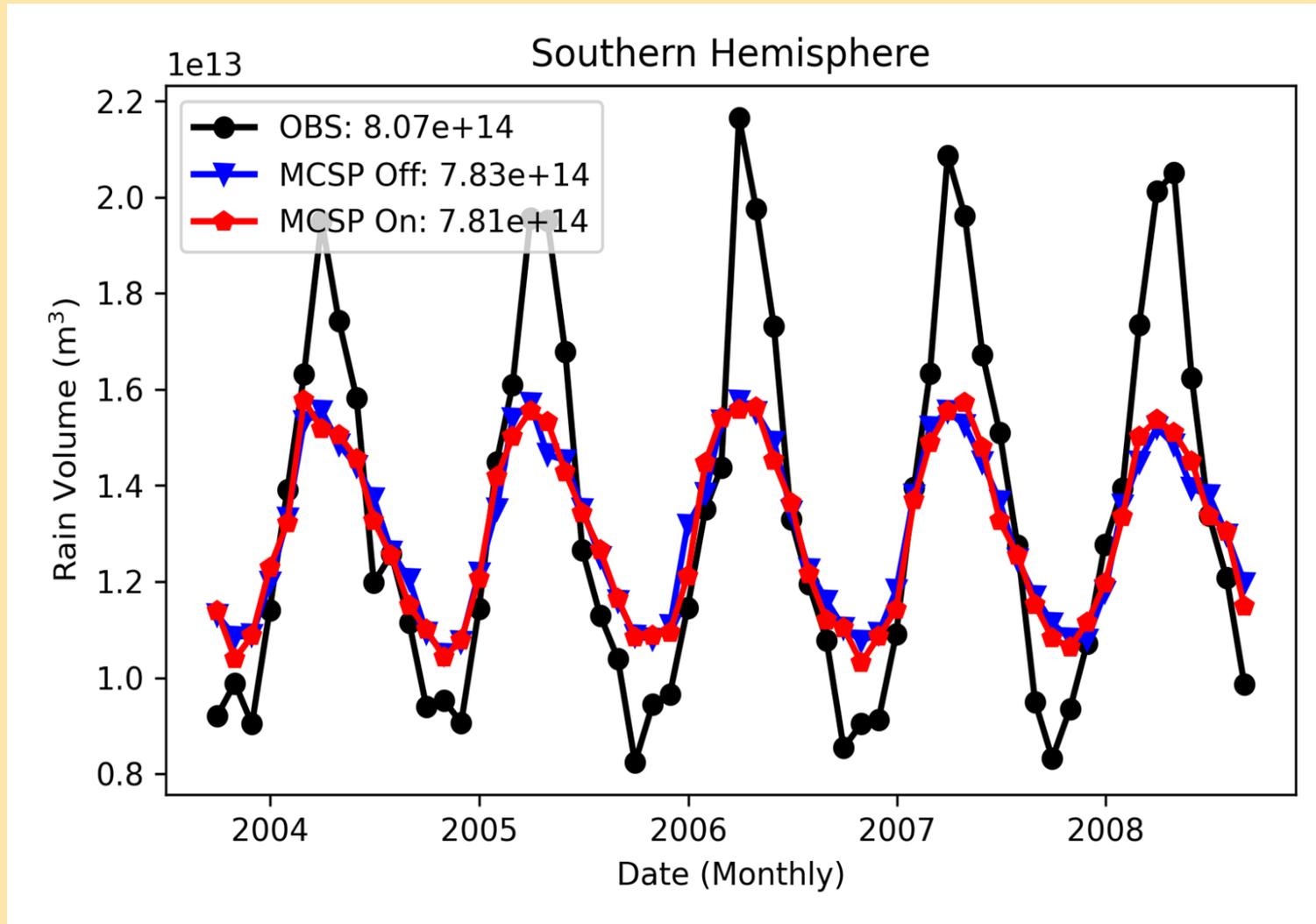
MCSP On - OBS  
correlation: 0.886



# Change in absolute bias

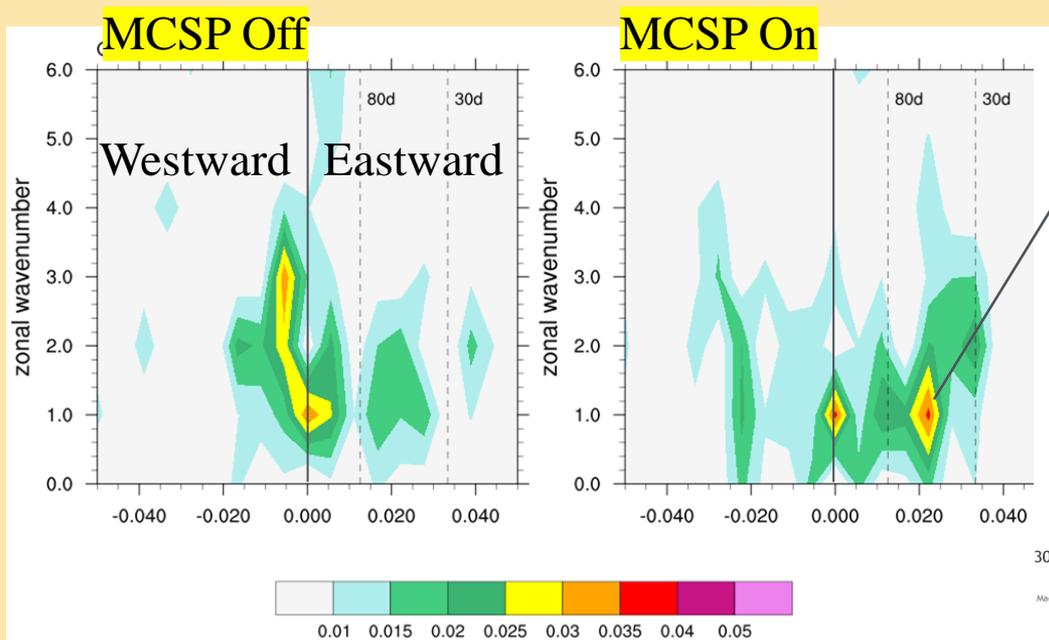


# Precipitation seasonal cycle

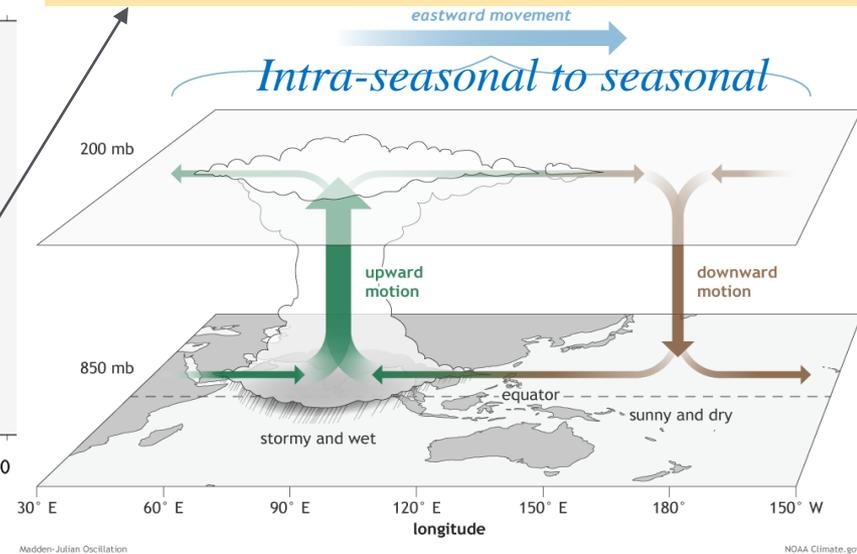


No change in S hemisphere

# Tropical wave frequency



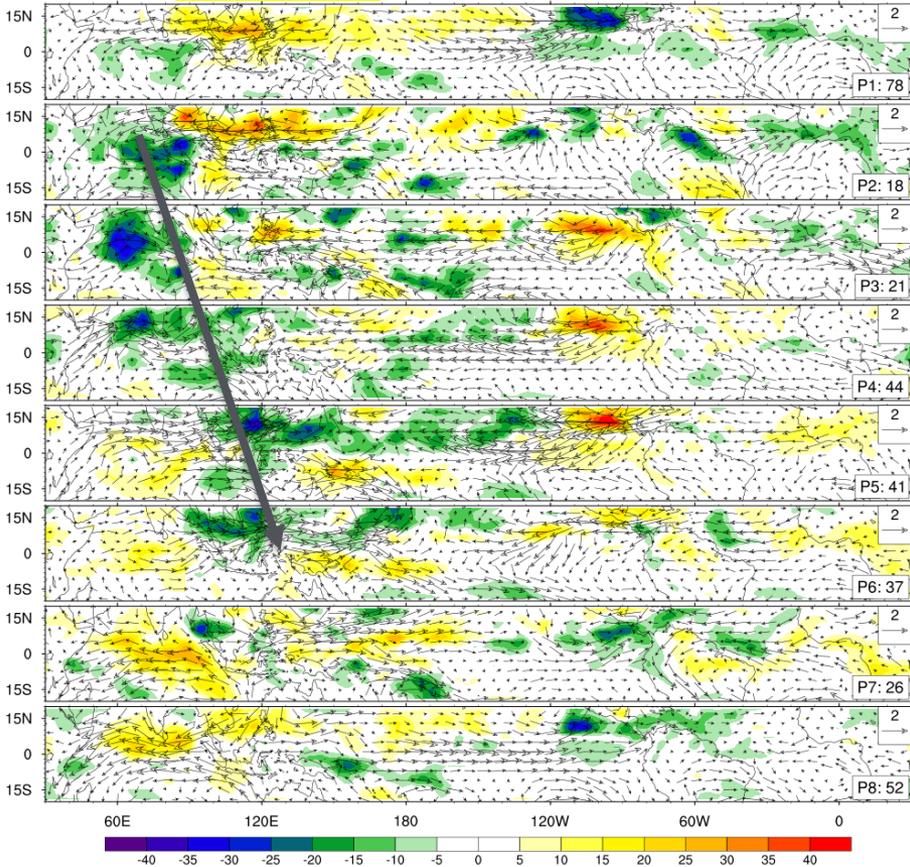
## Madden–Julian Oscillation (MJO)



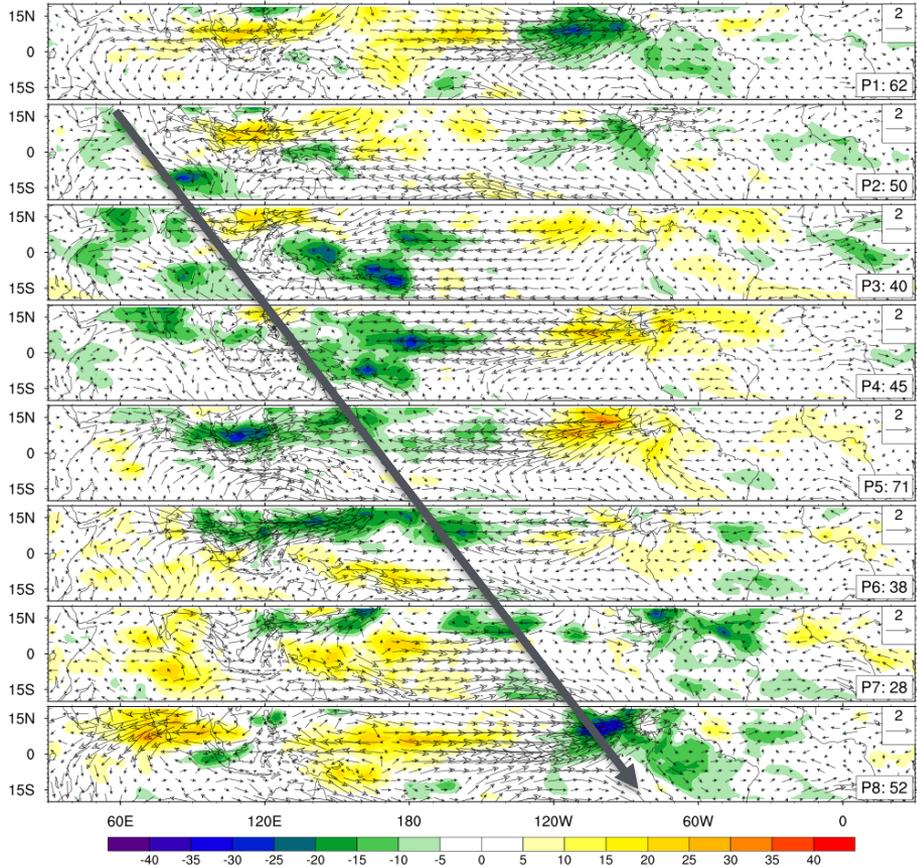
MCSP scheme enhances the plausible eastward-propagating waves and suppresses the unrealistic westward-propagating waves (May–Oct spectra)

# MJO life cycle, OLR and U850 anomalies

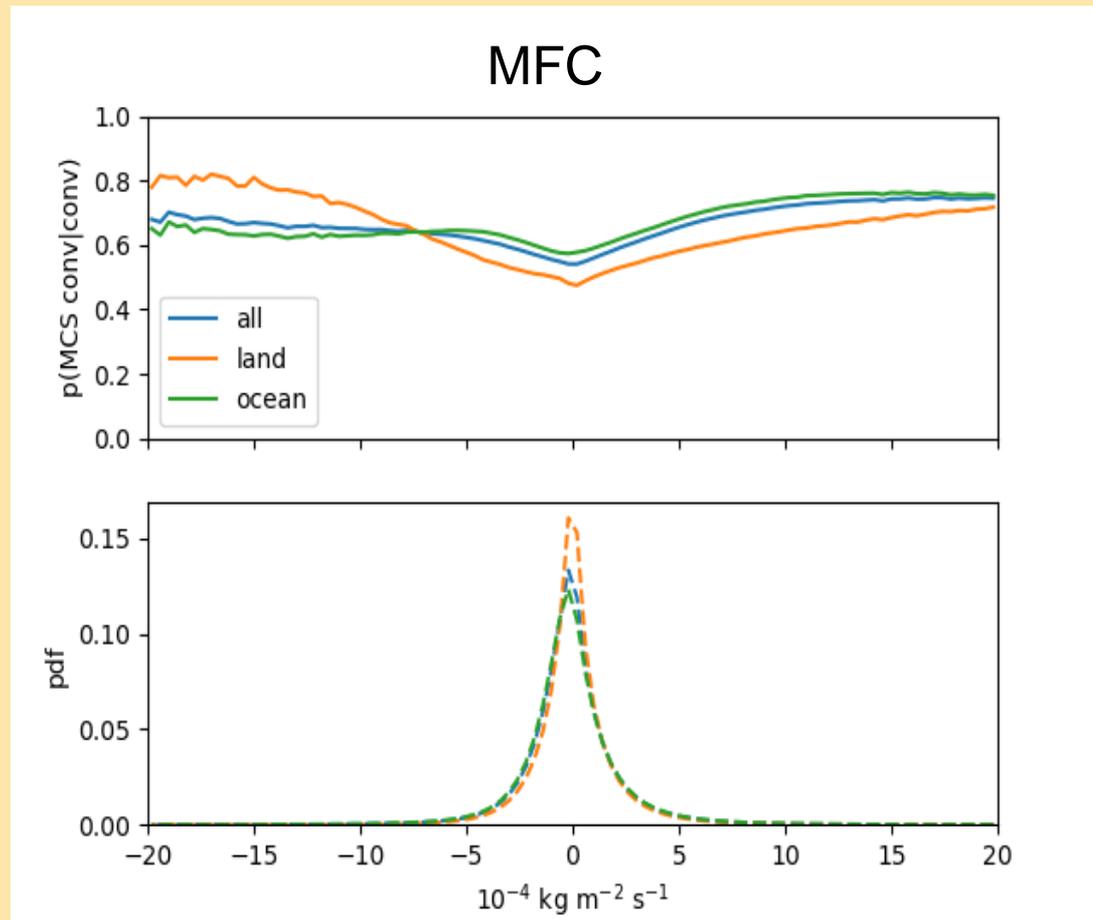
**MCSP Off**: 2004-2007: May to Oct



**MCSP On**: 2004-2007: May to Oct



# Scheme needs P (MCS | conv, variable)



Conditional  
probs

prob of  
variable

- Some change with MFC but TCWV is more marked