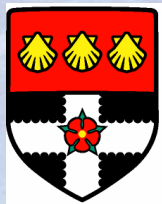


Friction in Mid-latitude Cyclones

Ian Boutle

Bob Plant, Stephen Belcher, Bob Beare, Andy Brown



**University of
Reading**



**National Centre for
Atmospheric Science**
NATURAL ENVIRONMENT RESEARCH COUNCIL



Motivation

- Many studies have shown the significance of friction in formation and dissipation of cyclones
- Up to 50% reduced growth rate
- Met Office Unified Model still has trouble with cyclone depth
- Underlying physical mechanism not well understood

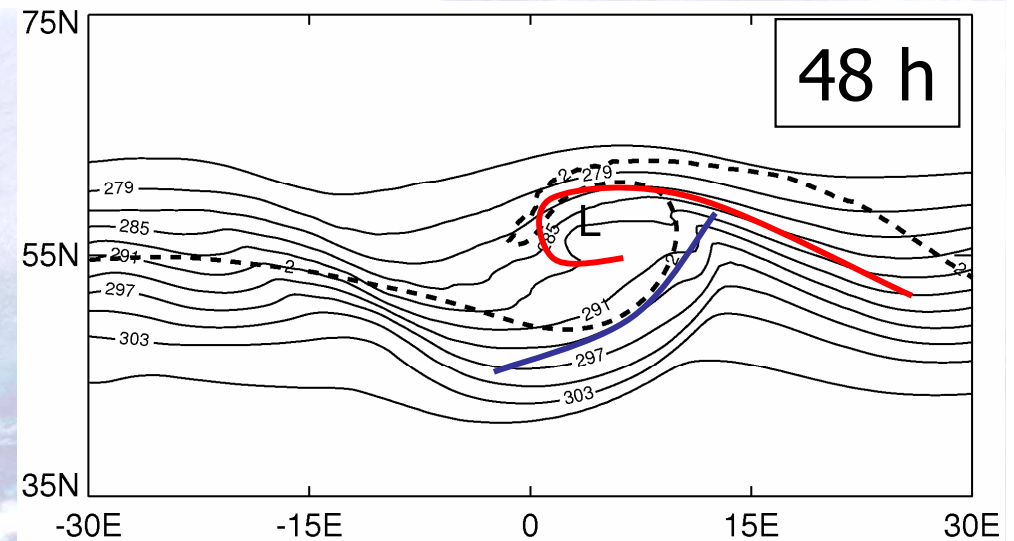
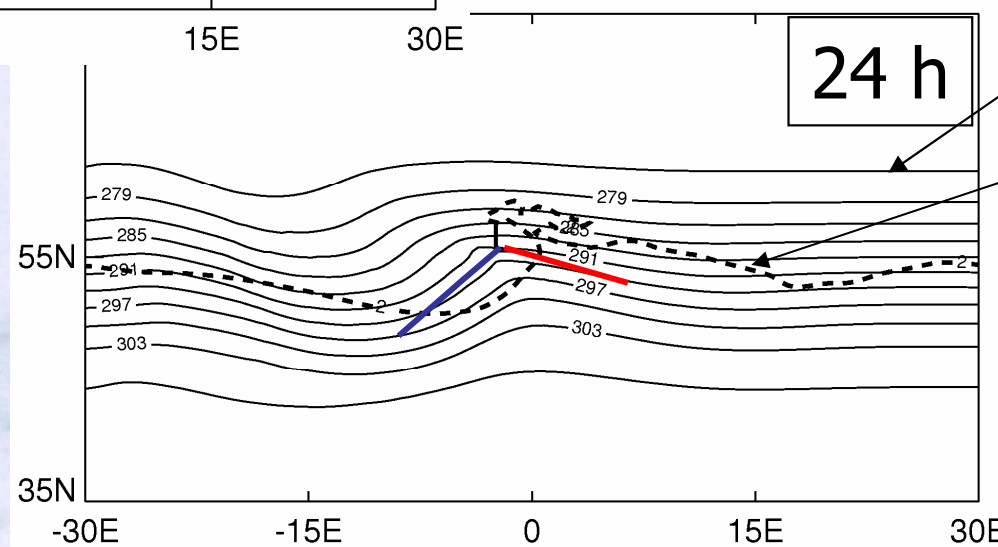
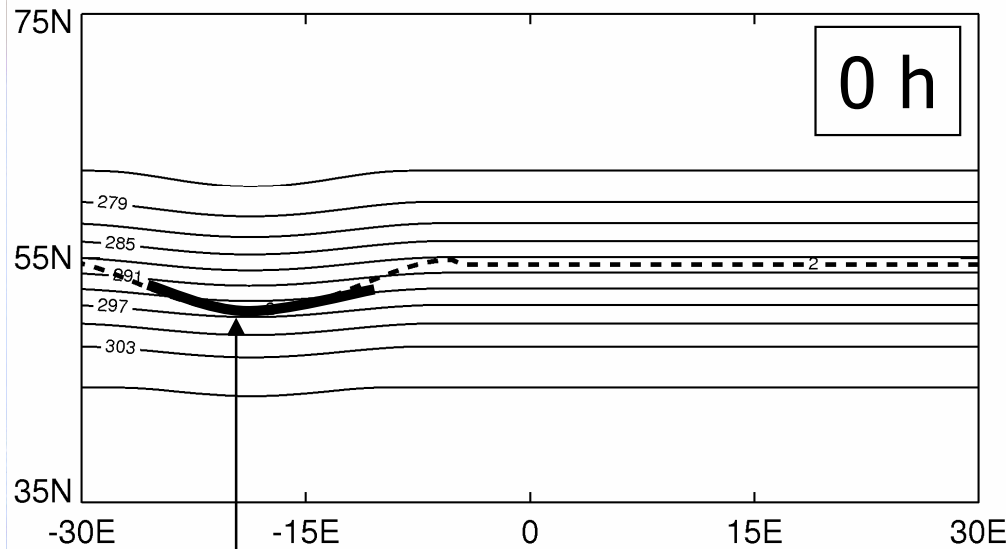
Model Setup

Near-surface Temperature

Tropopause

Upper Level Trough

Idealised, dry setup
– only dynamics
and boundary layer
scheme operating

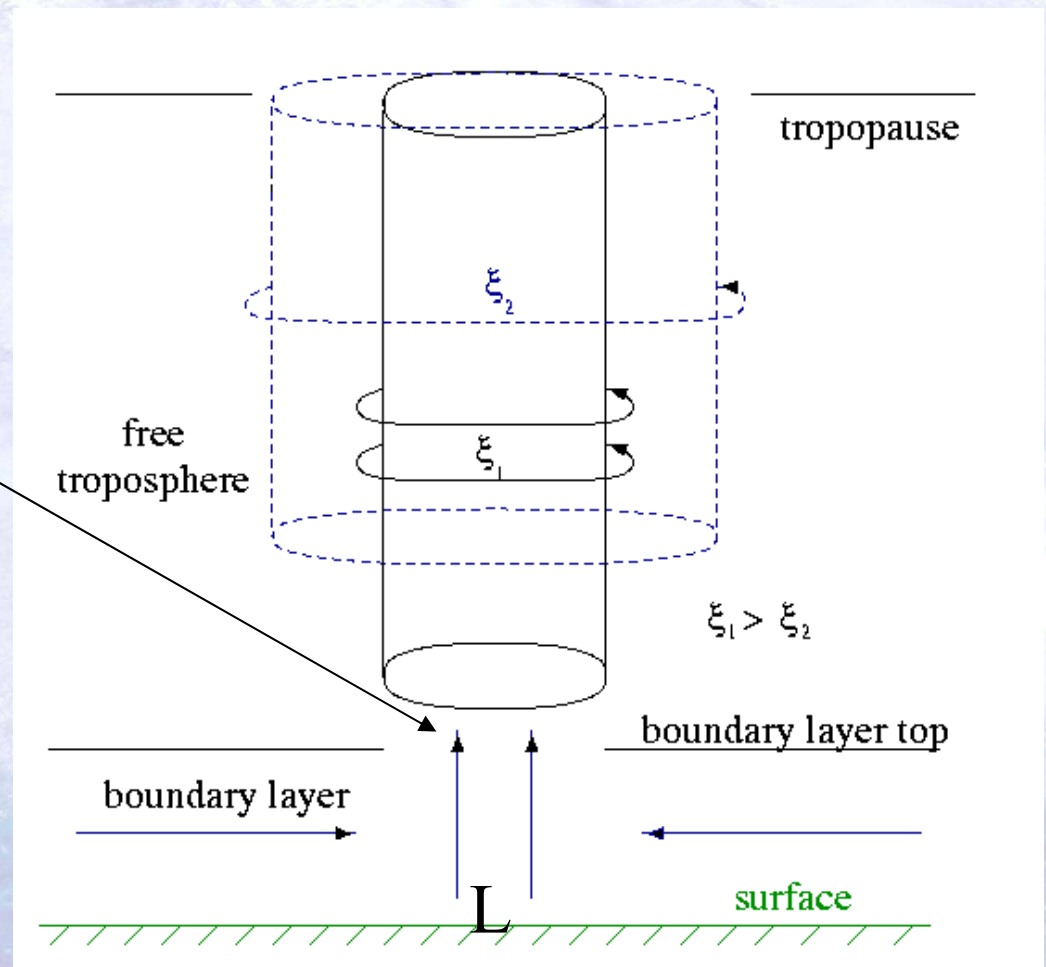


Ekman Pumping

- Boundary layer forces convergence
- Continuity forces ascent:

$$w_e = \frac{1}{f_0} \mathbf{k} \cdot \nabla \times \boldsymbol{\tau}_s$$

- Ascent squashes vortex, reducing vorticity
- What about temperature?



Potential Vorticity

$$PV = \frac{1}{\rho} \zeta \cdot \nabla \theta$$

$$\frac{D}{Dt}(PV) = \frac{1}{\rho} \left(\zeta \cdot \nabla \left(\frac{D\theta}{Dt} \right) + (\nabla \times \mathbf{F}) \cdot \nabla \theta \right)$$

Diabatic Term:

- Surface heat fluxes
- Latent heat fluxes

Frictional Term:

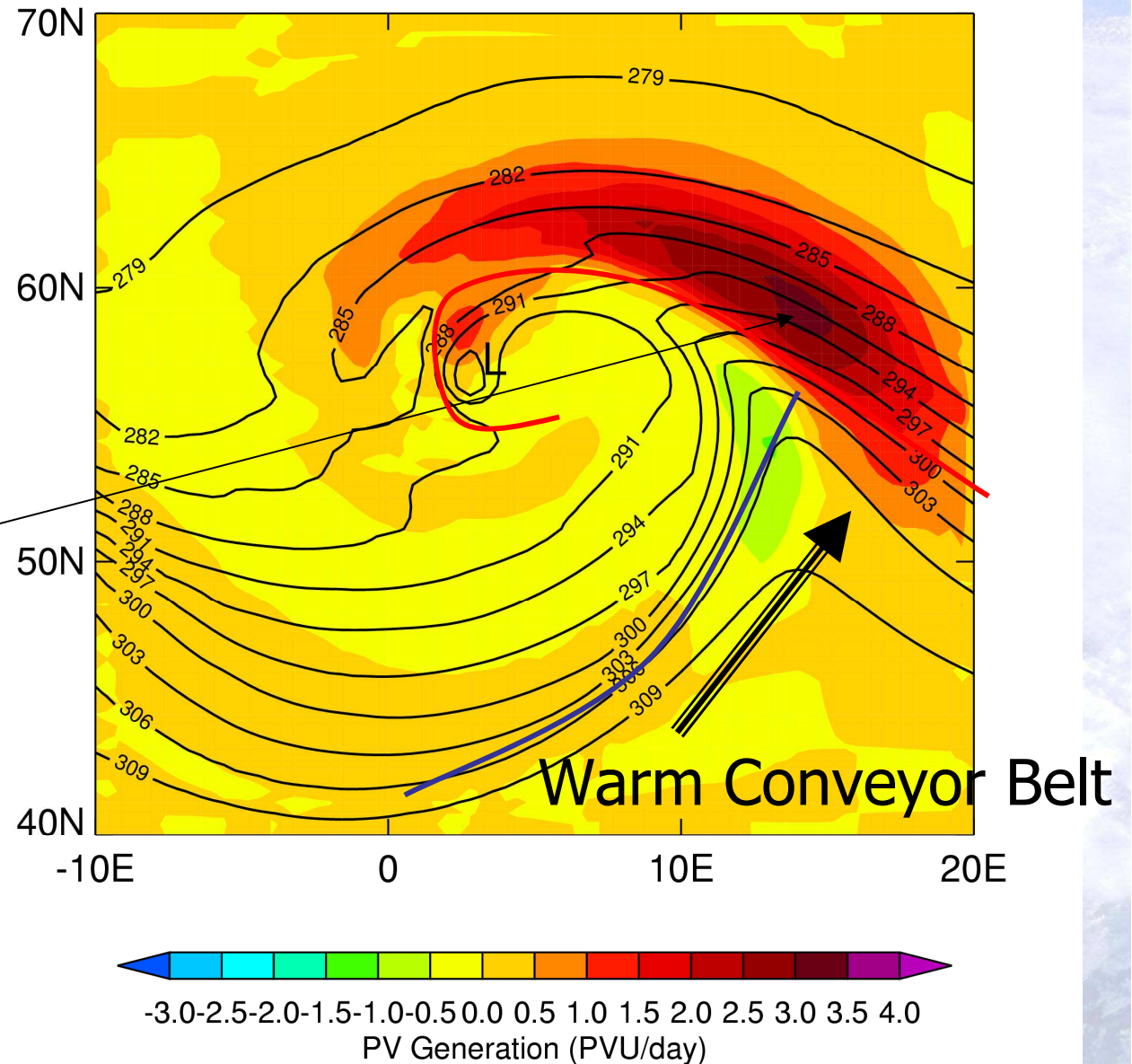
- Ekman pumping
- Baroclinic mechanism

Baroclinic Mechanism

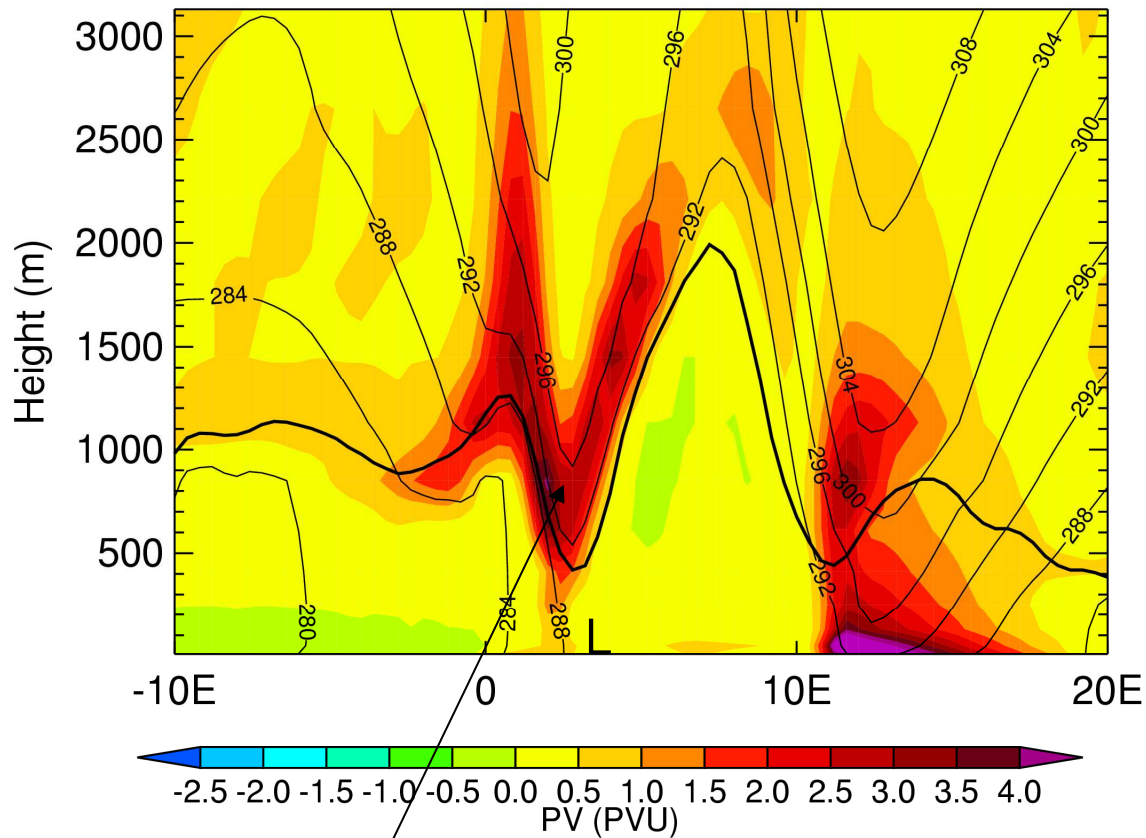
Depth averaged PV generation in boundary layer:

$$[G_B] = \frac{1}{\rho^2 h^2} \mathbf{k} \times \boldsymbol{\tau}_s \cdot (\nabla_H \theta)_h$$

Depends on alignment of surface and thermal winds

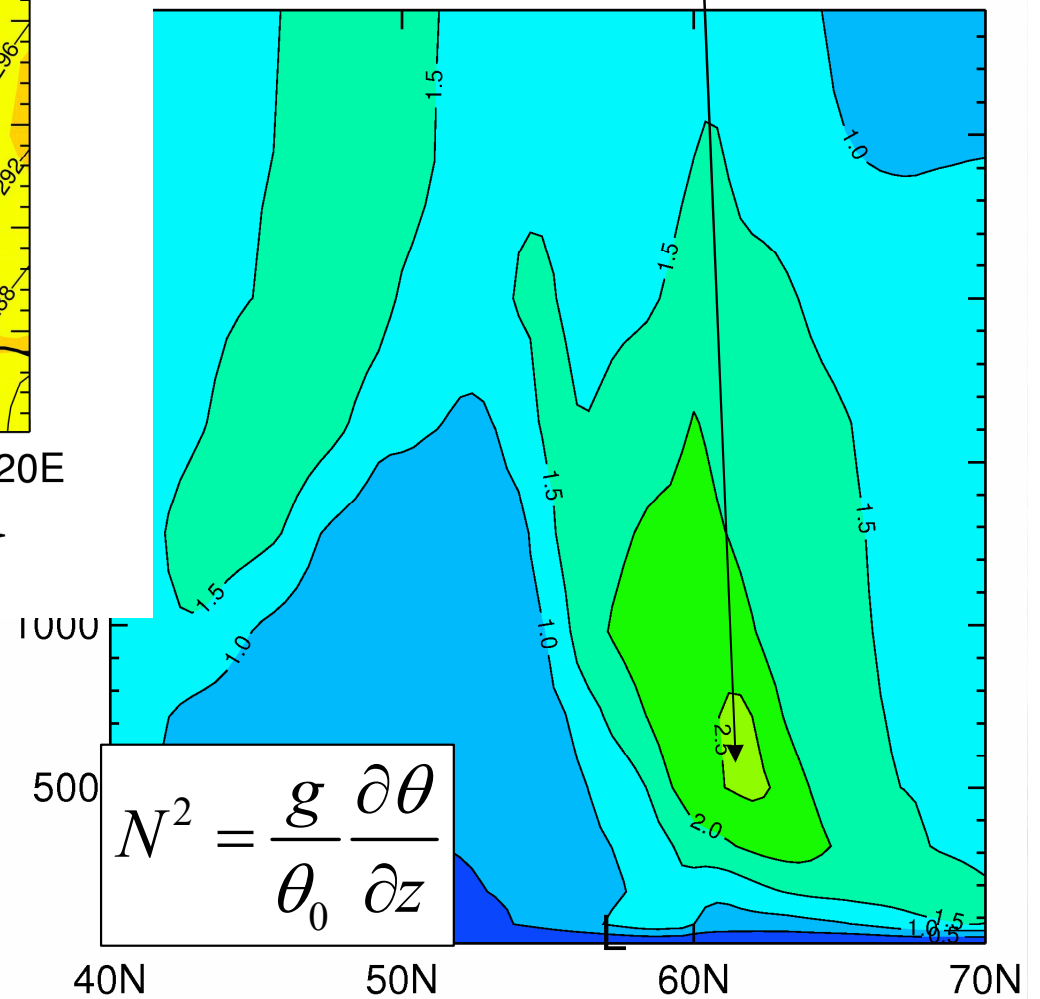


Baroclinic Mechanism 2

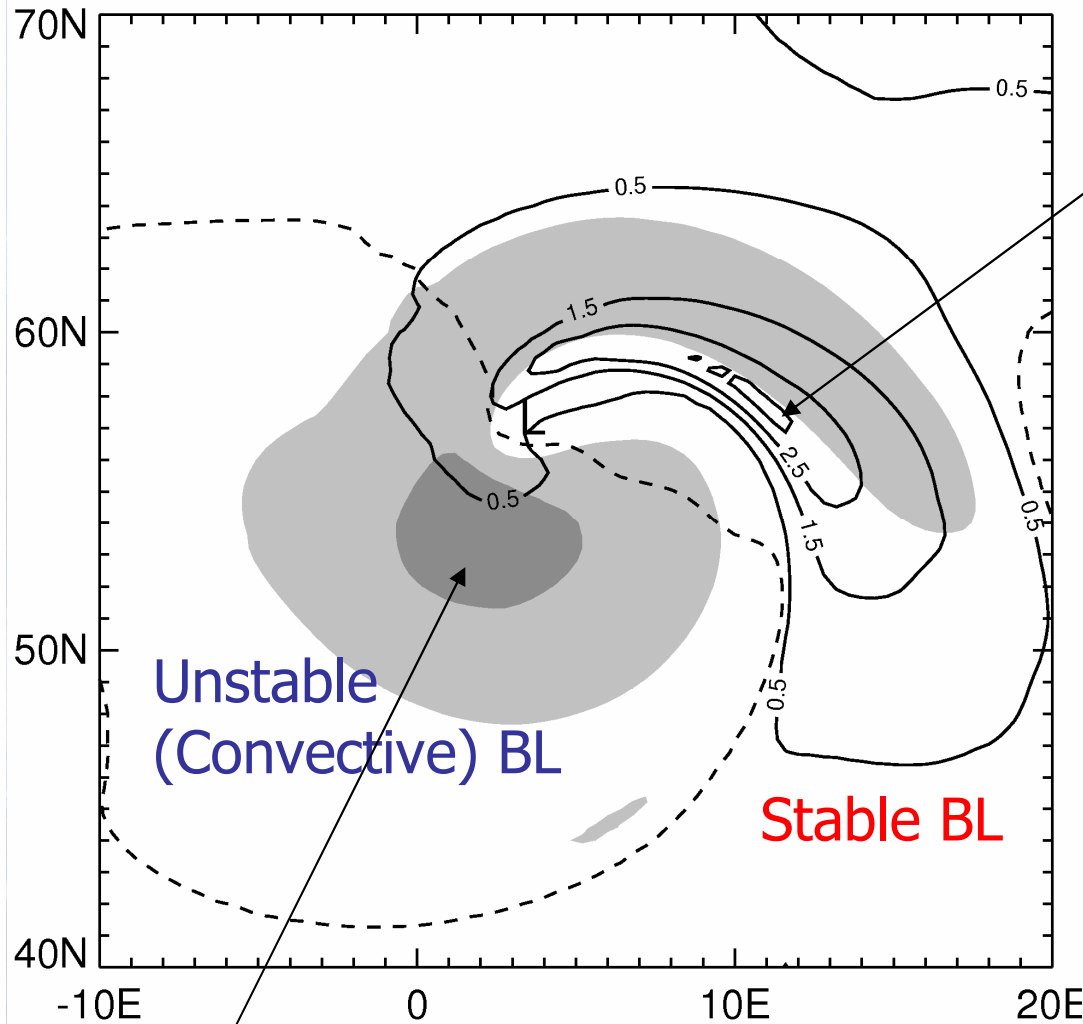


PV accumulates above
Low centre

Mainly Static
Stability Anomaly



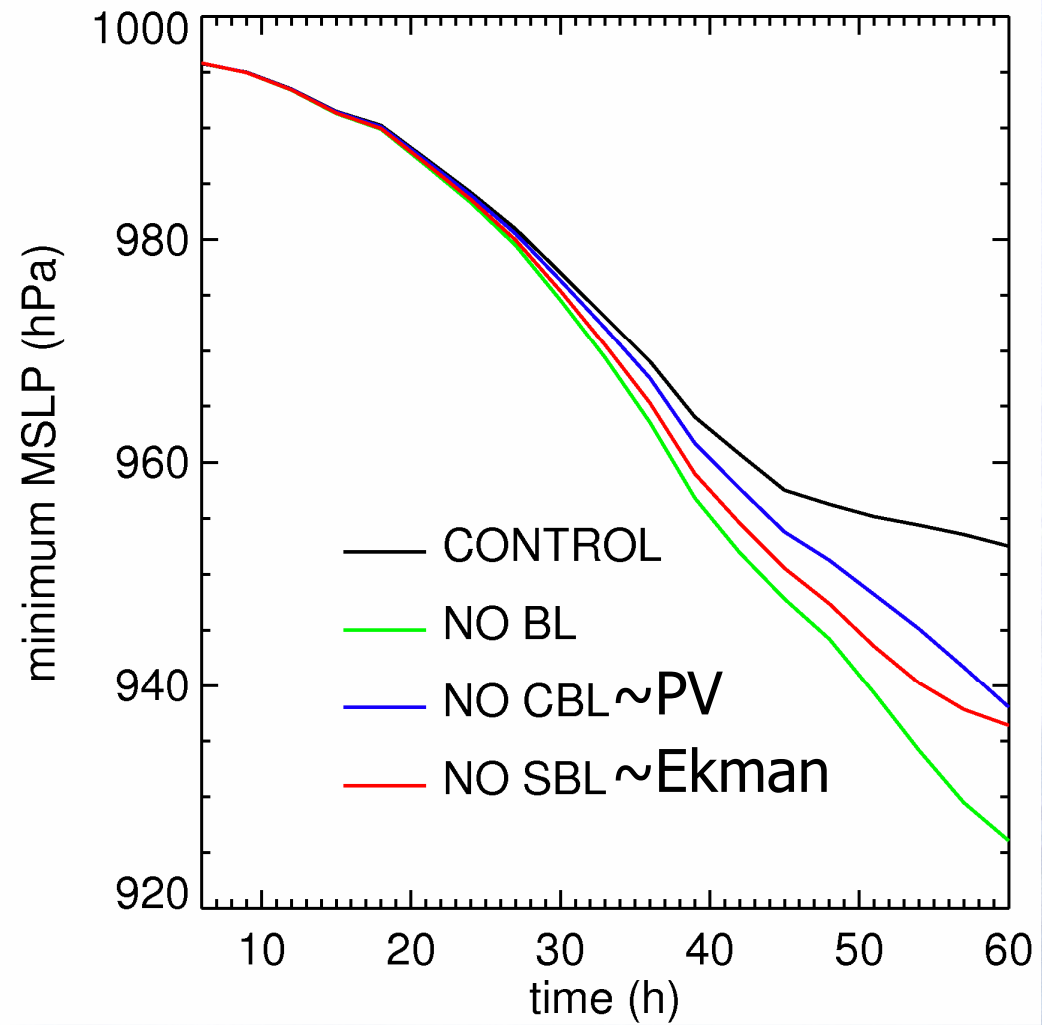
Boundary Layer Stability



PV mainly
in SBL

Unstable
(Convective) BL

Stable BL



Surface
stress
mainly
In CBL

Summary

- Boundary layer friction has a large affect on cyclone development
- Ekman pumping is significant, but not the only process acting
- PV perspective provides another mechanism, which appears similarly important

Switching off BL

- Surface stress parameterised in terms of “eddy diffusivities”

$$\tau_{xz} = \rho K_m \frac{\partial u}{\partial z}$$

- Define SBL by $\theta(\text{level } 1) > \theta(\text{surface})$
- Set $K_m = 0$ here