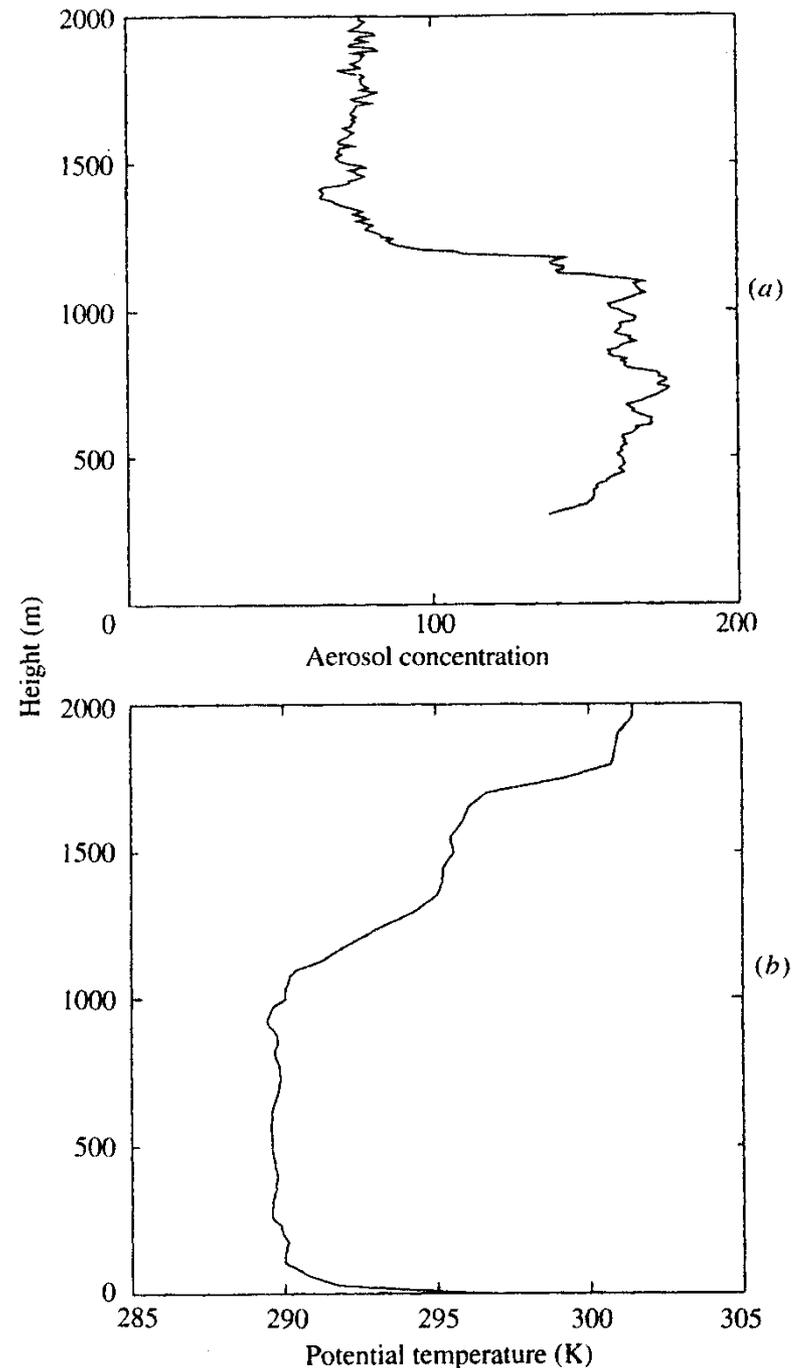


Boundary layer ventilation by weather systems

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Motivation

- Most emissions at surface
- Easily mixed through depth of boundary layer
- BL inversion a strong barrier to further vertical mixing
- Surface concentrations determined by:
 - Source strength
 - Boundary layer depth
 - Rate of ventilation of boundary layer
- Case studies
 - Ventilation rate varies
 - 20% - 70% mass ventilated

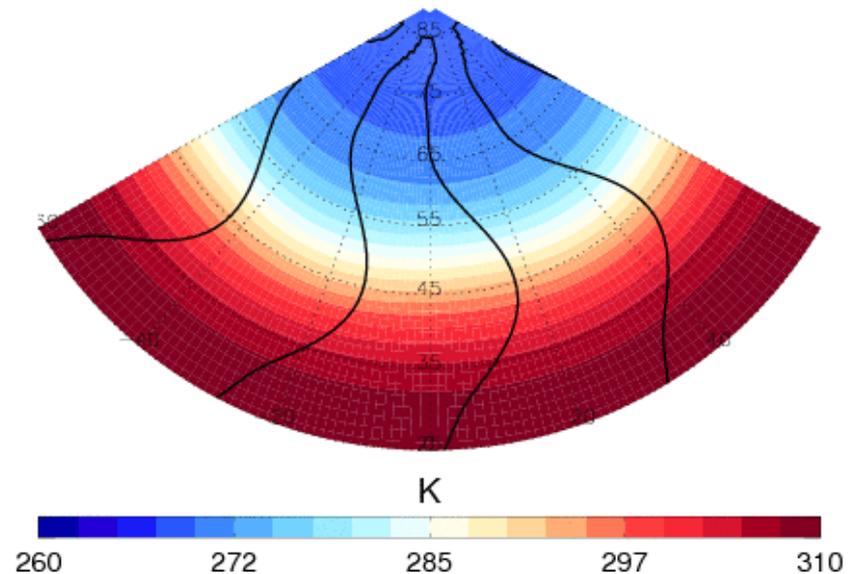


Idealised simulations

- Simulate idealised weather systems
- Use Met Office Unified Model in idealised channel configuration or IGCM
- **Dry simulations**
- Include Boundary-layer scheme
- Constant strength source at surface, across the whole domain

$$\frac{DC}{Dt} = \nabla \cdot \vec{F}$$

$$\vec{F} = F_s \quad \text{at } z = 0$$

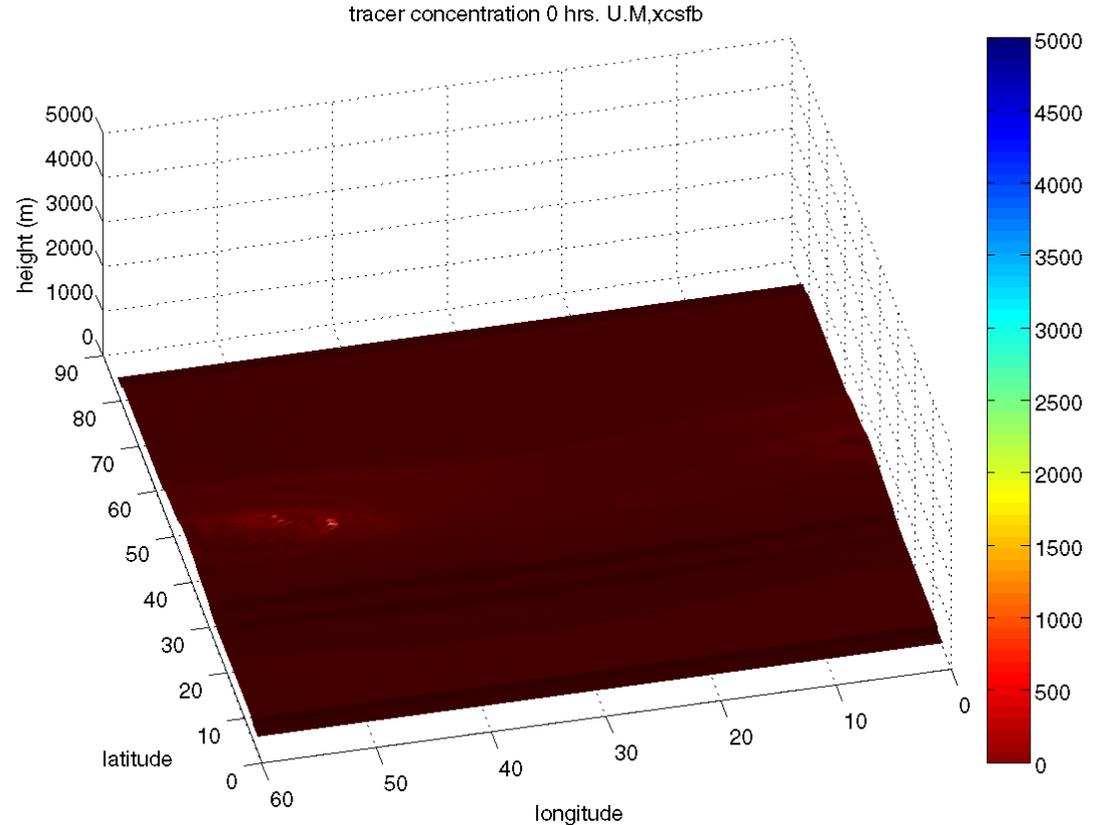


Lines: Surface pressure
Contours: potential temperature at 1km
Each frame is 1 day apart

Transport into free troposphere

Three phases:

- Vertical BL mixing
- Horizontal transport within BL
- Ventilation by warm conveyor belt



Isosurface of tracer concentration.
Colours show the height of the surface

What controls ventilation rate

- The amount of turbulent mixing within the boundary layer?
 - Pollutants need to be mixed up to near the top of the boundary layer for ventilation to occur
- Horizontal transport within the boundary layer?
 - Convergence and divergence within B.L.
 - Only certain regions of the boundary layer are ventilated
- The large scale vertical motion associated with the cyclone?
 - Final step in ventilation
 - Most important?
- Numerical experiments:
 - Role of boundary layer convergence: vary drag coefficient
 - Role of large scale flow: vary weather system strength

BL convergence

- Cyclone energy sensitive to BL drag
- Ventilation mass flux is not

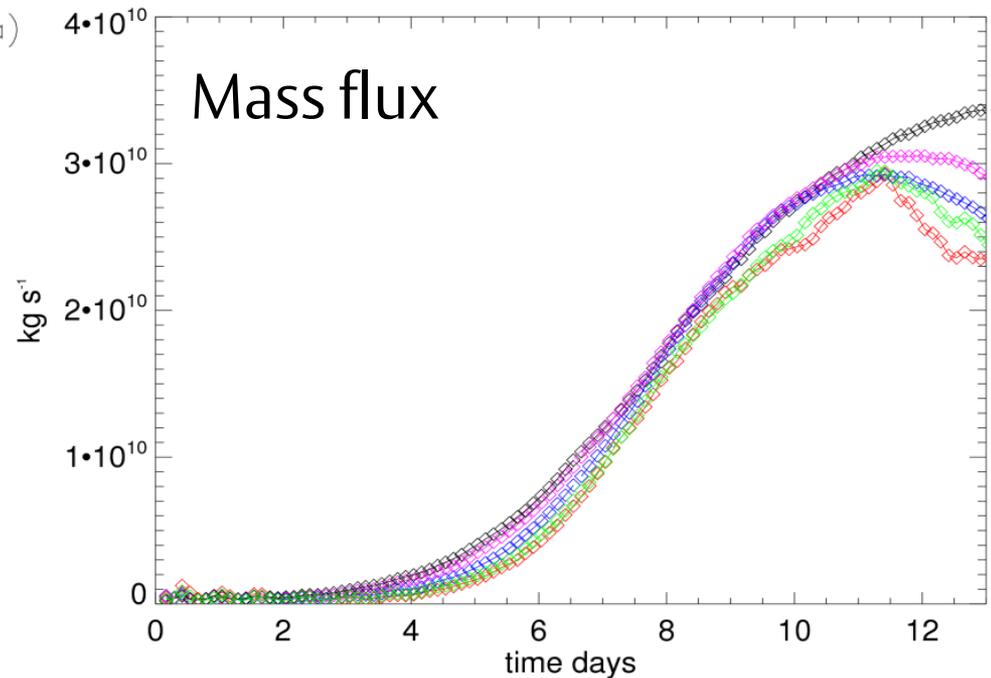
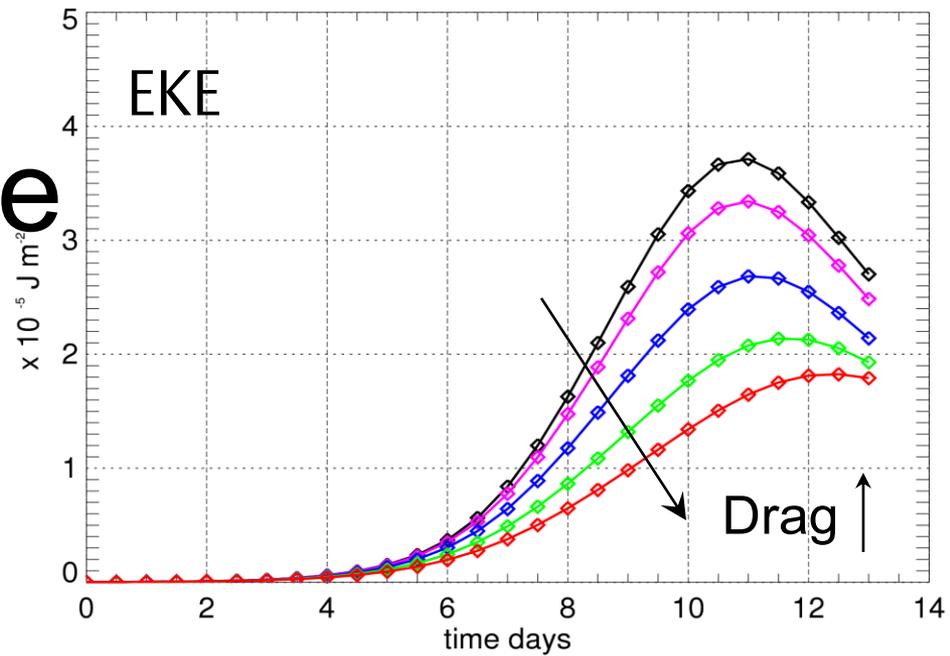
$$Cd = 0$$

$$Cd = 0.1 \times 10^{-3}$$

$$Cd = 0.56 \times 10^{-3}$$

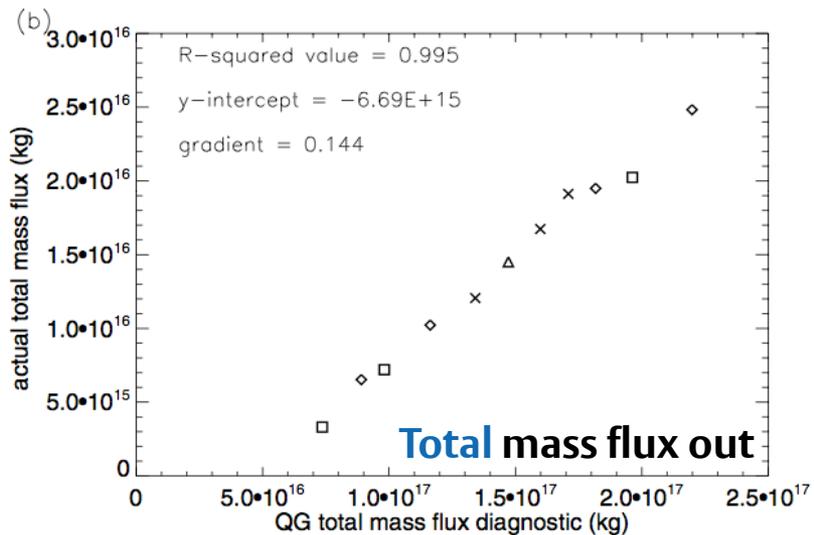
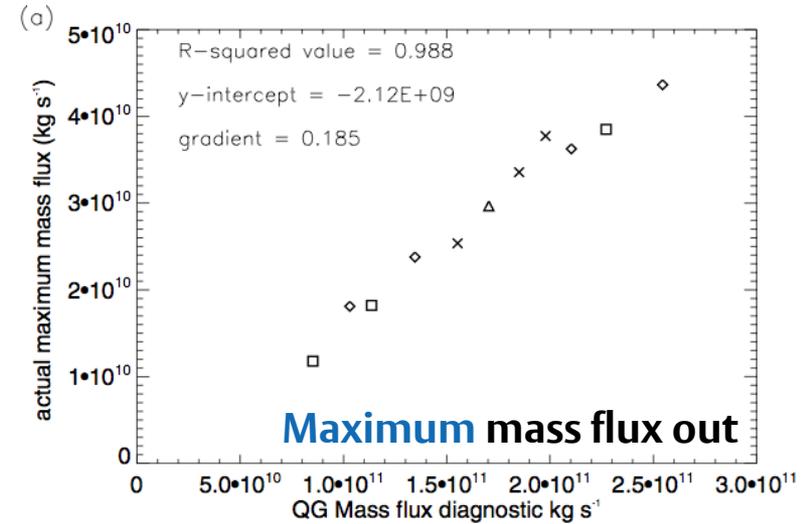
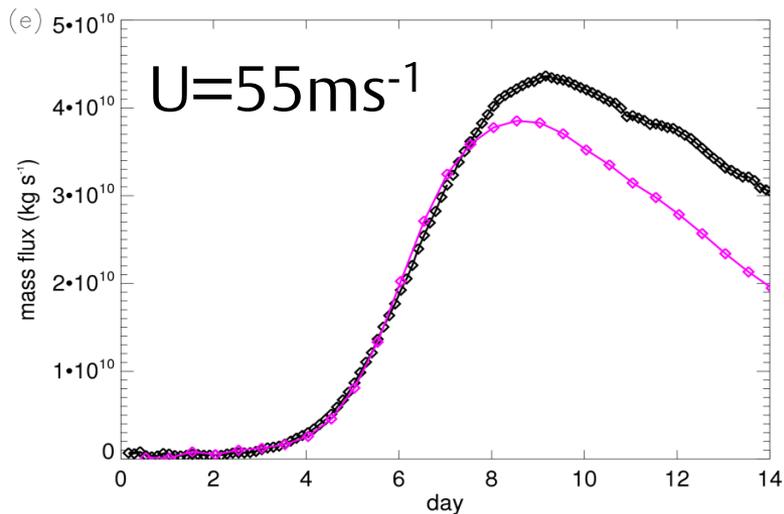
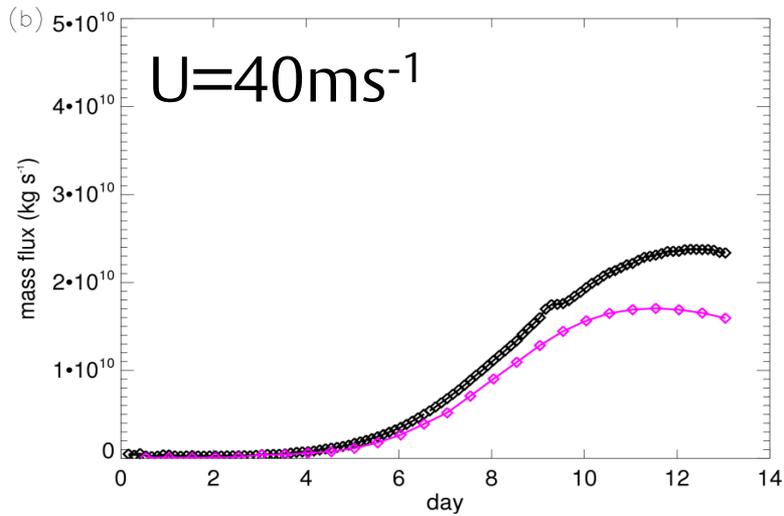
$$Cd = 2 \times 10^{-3}$$

$$Cd = 5 \times 10^{-3}$$



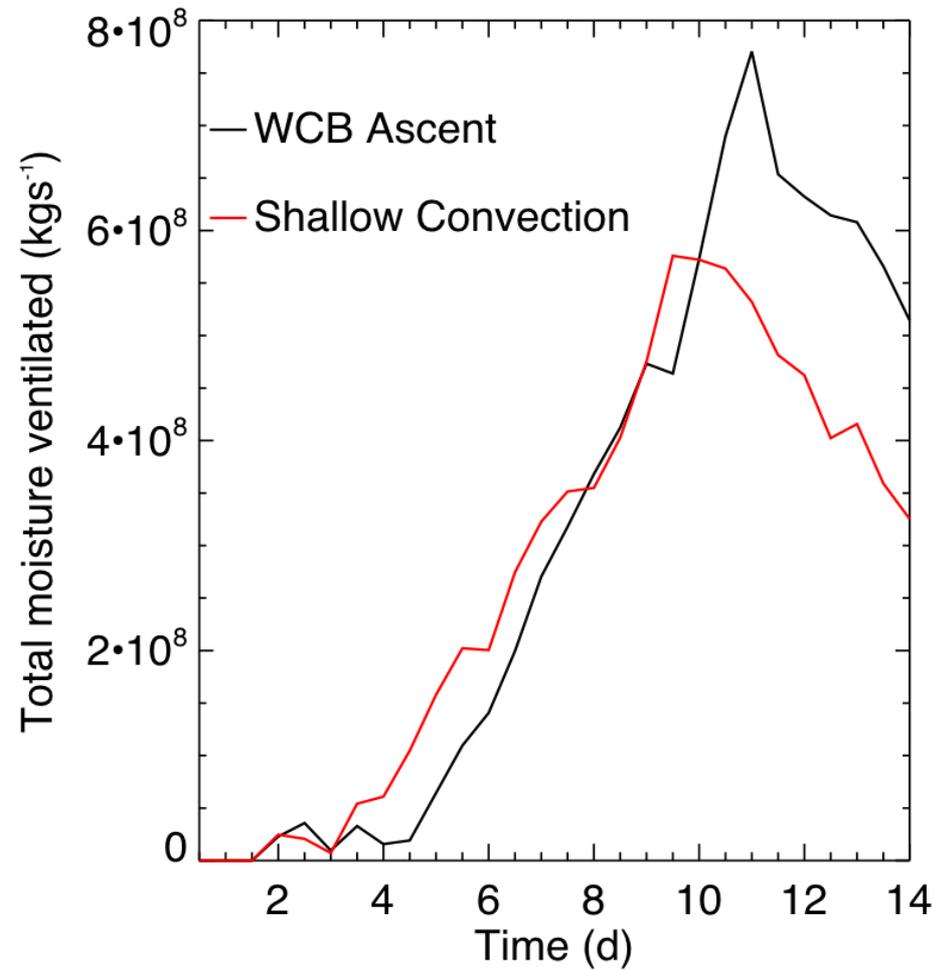
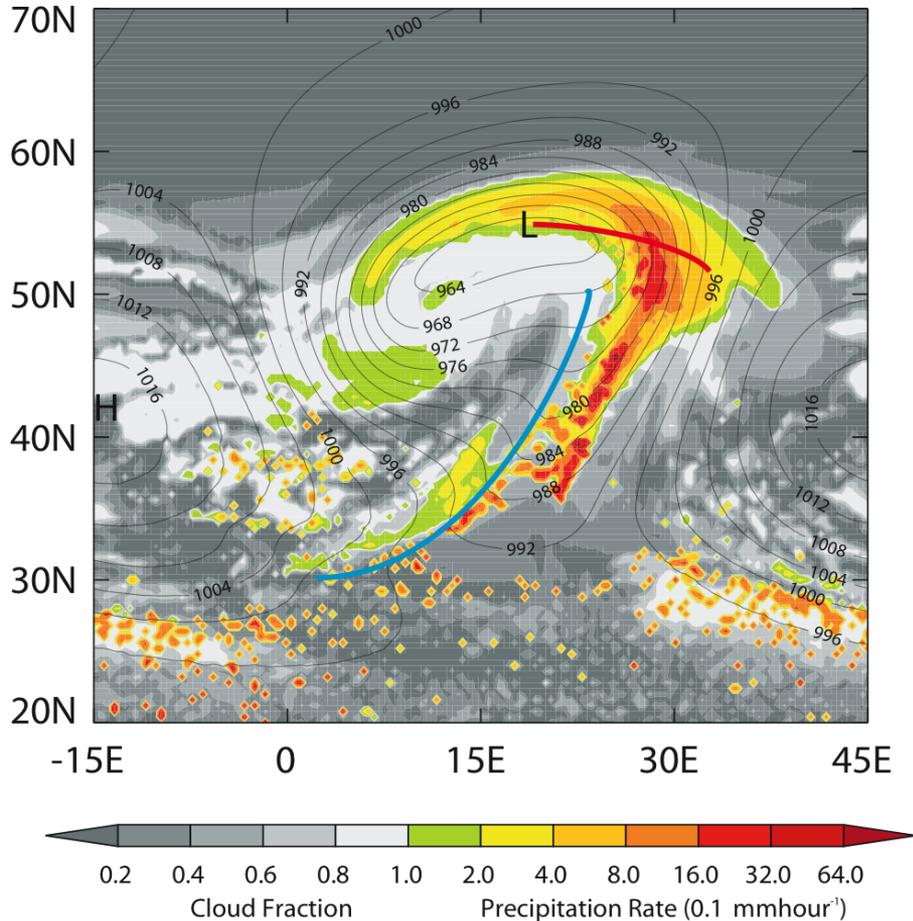
Role of large scale flow

$$w \propto \frac{-v_g f \frac{\partial u_g}{\partial z}}{2N^2}$$



Moist lifecycles

(b)



The sting:

moist convection as effective
as large scale ventilation

Conclusions

- BL top acts as barrier to vertical transport
 - Comparable to strat-trop exchange
- Mid-latitude weather systems ventilate boundary layer
 - Large-scale vertical WCB motion is the dry control
 - Re-analysis products sufficient to capture WCB
- Moist lifecycles:
 - New pathway through convection
 - Convection gives comparable ventilation as WCB
 - More delicate to represent in models

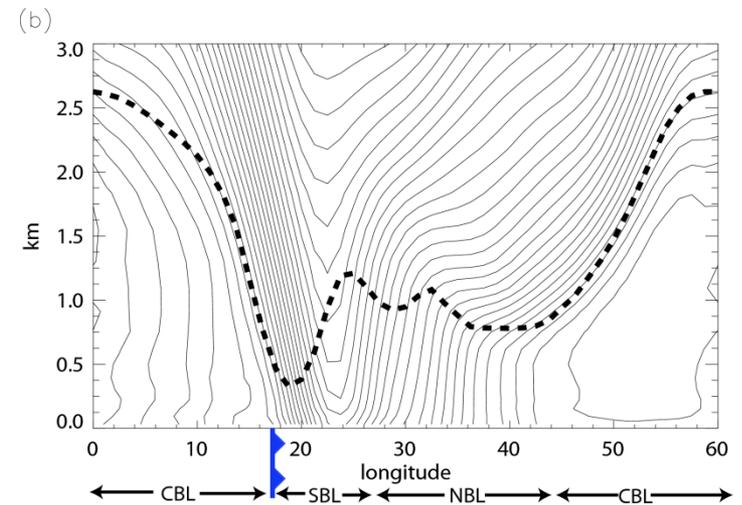
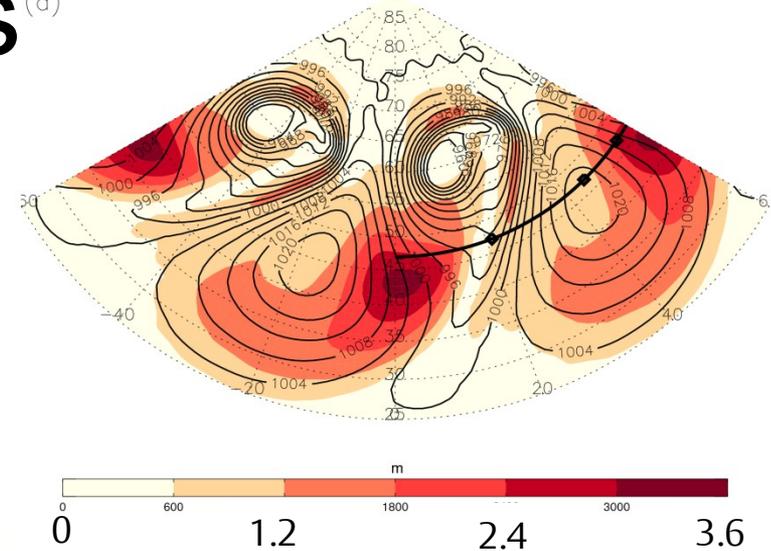
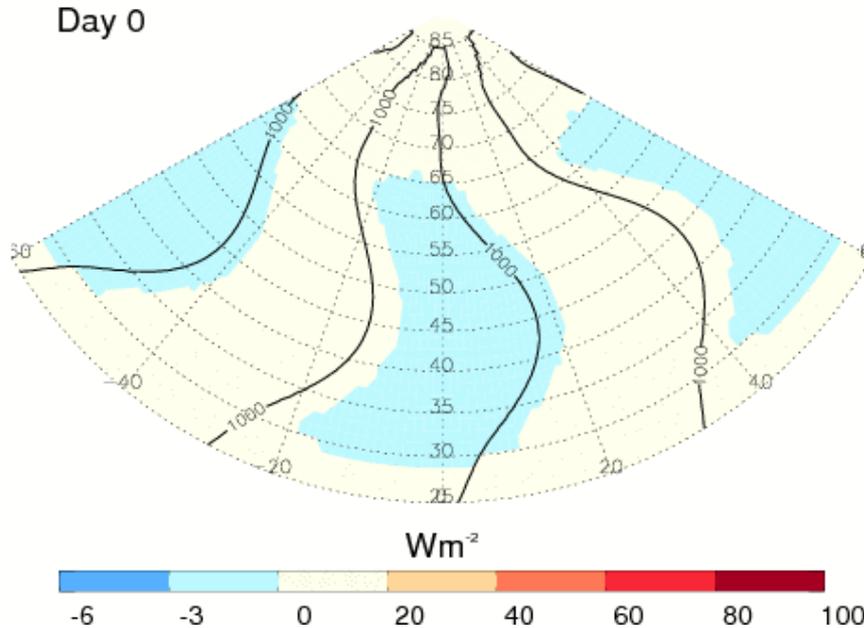
Conclusions

- Pathways for boundary layer ventilation:
 - Diurnal cycle of BL
 - Transport through inhomogeneous BL
 - Convection:
 - case specific
 - Fast: 3 hours up to 5km
 - Conveyor belts:
 - Reliable: all cyclones
 - Slow: 12 hours to 5km

Further questions

- How does the chemistry care:
 - Time scale for ascent?
 - Height of destination?
 - Passage through clouds?
 - When do trajectories fail?

B.L. characteristics^(a)



- Fluxes drive by thermal advection
- Deepest boundary layers are co-located with maximum heat fluxes
- Large change in boundary layer depth across cold front

Boundary layer mass budget

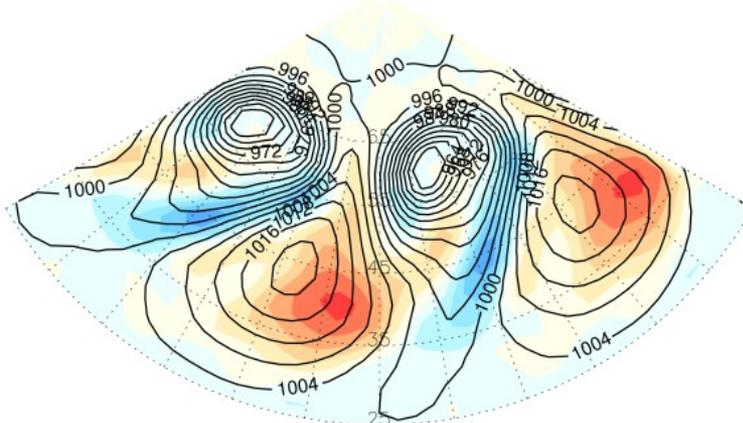
- Integrate continuity over the B.L depth:

$$\frac{\partial \tilde{\rho}}{\partial t} = \rho_h \frac{\partial h}{\partial t} + \overbrace{\text{Convergence}}^{\text{Synoptic}} - \rho_h (\underline{u} \cdot \underline{n})_h$$

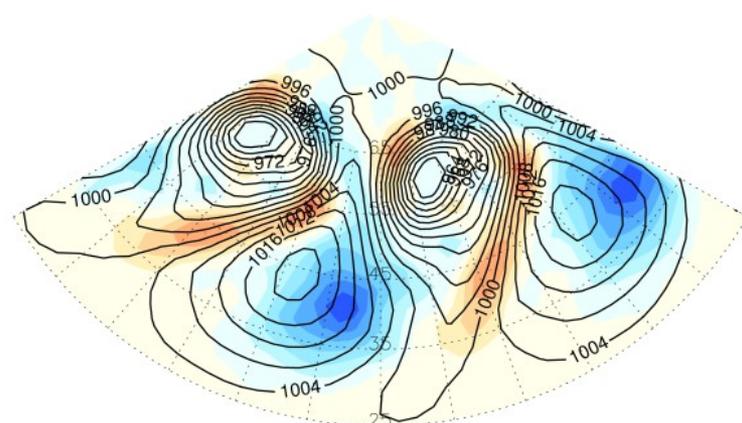
↑
↑
↑
 Changes in B.L depth Transport within B.L. Synoptic ascent/descent

Synoptic

Convergence



kg s⁻¹



kg s⁻¹



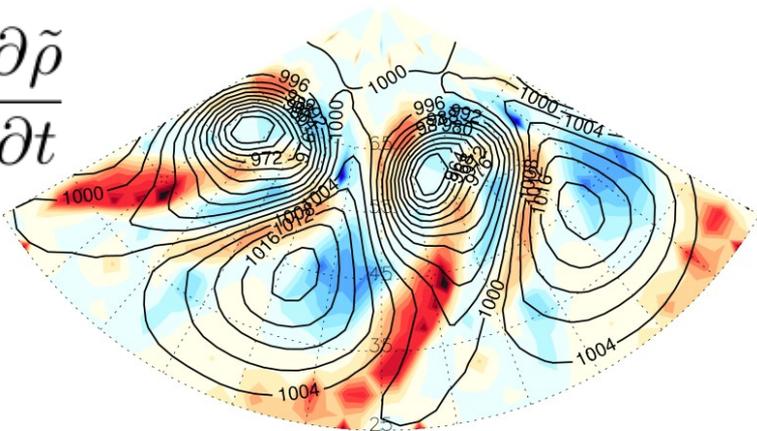
Boundary layer mass budget

- But B.L depth can change due to the **surface heat flux**, which leads to entrainment into the boundary layer
- OR due to the **large-scale vertical motion (subsidence)** pushing down the theta contours
- Combine to form an 'entrainment' velocity

$$w_e = \frac{\partial h}{\partial t} - (\vec{u} \cdot \vec{n})_h$$

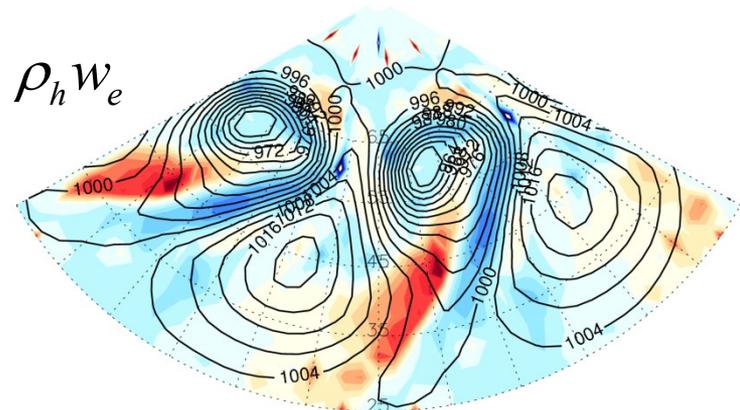
$$\frac{\partial \tilde{\rho}}{\partial t} = \rho_h w_e + \tilde{C}_h$$

$\frac{\partial \tilde{\rho}}{\partial t}$



kg s⁻¹

$\rho_h w_e$

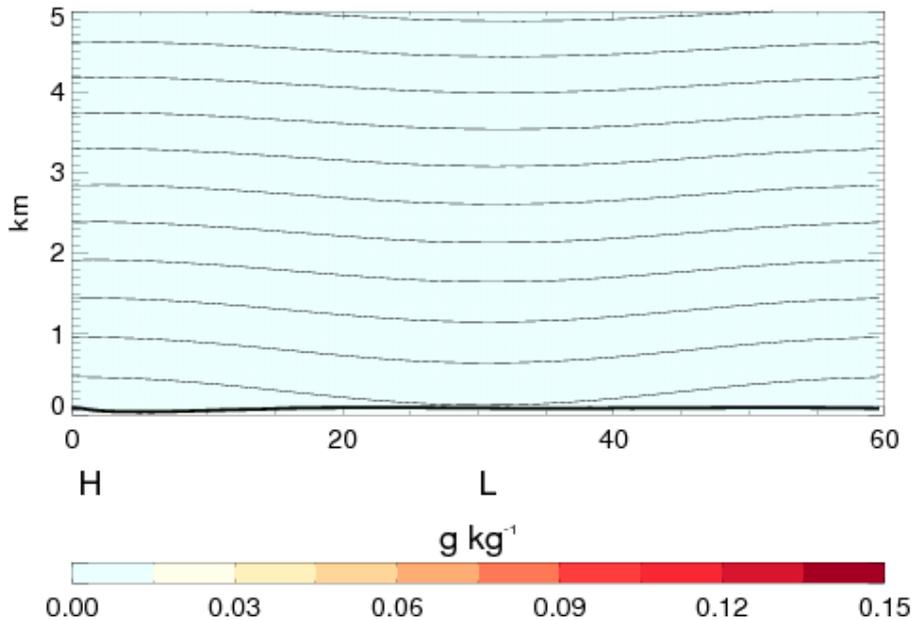


kg m⁻² s⁻¹

-3.0e+09 -1.5e+09 0.0e+00 1.5e+09 3.0e+09

-5.0e-02 -3.0e-02 -1.0e-02 1.0e-02 3.0e-02 5.0e-02

Transport of tracer in the B.L.



X-Z cross section. Potential temperature (contours), tracer concentration (colours) and B.L. depth (black line)

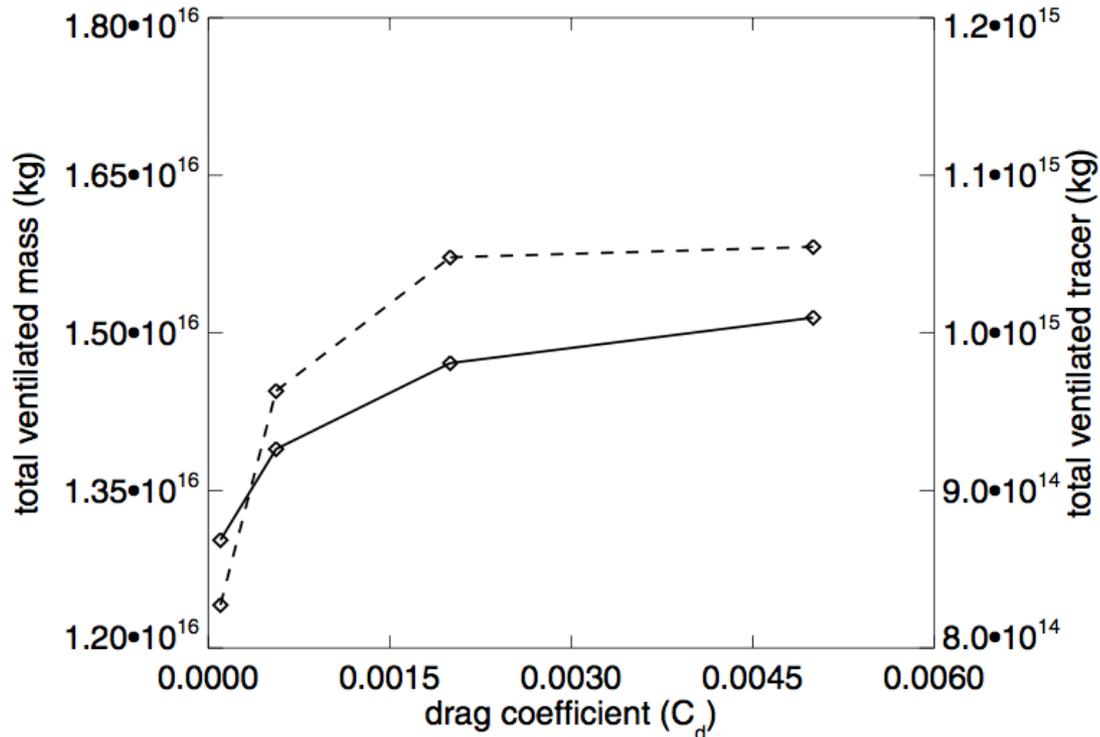
Tracer is mixed vertically by turbulent mixing .

- More mixing in high pressure region
- Little vertical mixing in low pressure region

Tracer is also transported horizontally within the B.L.

- Convergence towards low centre
- Divergence out of high pressure

Mass ventilated



- Surface type has little affect on mass ventilated
- When 'no' friction acts, reduced mass ventilated
- No tracer in the correct regions
 - conveyor belt footprint areas or the mid to upper regions of the B.L