

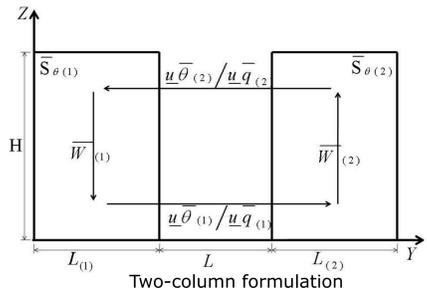
# Modelling the interactions between tropical convection and large scale dynamics

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## 1. Introduction

Radiative-convective equilibrium (RCE) simulations hide information on how large-scale dynamics feedback on convection and vice versa. To allow such feedbacks to occur, the **weak temperature gradient** (WTG) approximation of Sobel and Bretherton (2000) is used to couple a cloud resolving model (CRM) to a reference column (reference column model). In such a model, heat and moisture budgets are not strictly closed. A large-scale circulation develops over a homogeneous environment (results similar to those of Raymond and Zeng (2005) and Perez et al. (2006)). In addition, the model has a **unique final state** no matter how it is initialised. The WTG approximation is again used to coupled two columns of the CRM (coupled columns model). The budgets are now closed. Such setup is more **realistic** and hasn't been studied before. The two models are compared. A large-scale circulation is not maintained in this model over a homogeneous environment. Hence, the large-scale circulation which develops in the reference column model is **artificial**.

## 2. WTG calculations



Gravity waves time scale. It is related to the length scale of large-scale circulation  $\chi = (L_{(1)} + L_{(2)}) / 2 + L$

Large-scale vertical velocities:

$$\bar{\omega}_{(1)} = -\frac{\bar{\theta}_{(1)} - \bar{\theta}_{(2)}}{\tau \left( \frac{\partial \bar{\theta}_{(1)}}{\partial z} + \frac{\partial \bar{\theta}_{(2)}}{\partial z} \right)}, \quad \bar{\omega}_{(2)} = -\bar{\omega}_{(1)} \quad \text{with } L_{(1)} = L_{(2)}$$

The tendency of  $\bar{\theta}_{(1)}$  due to large-scale circulation.

$$\frac{\partial \bar{\theta}_{(1)}}{\partial t} = \bar{\theta}_{(1)} \frac{\partial \bar{\omega}_{(1)}}{\partial z} - \frac{\partial \bar{\omega}_{(1)} \bar{\theta}_{(1)}}{\partial z}. \quad \text{Similarly for } \bar{q}_{(1)}, \bar{\theta}_{(2)} \text{ and } \bar{q}_{(2)}.$$

Heat and moisture budgets:

$$\sum_{i=1,2} \left\{ SHF_{(i)} + C_p \int_{surf}^{z_{top}} \rho \left[ \left( \frac{\partial \bar{T}_{(i)}}{\partial t} \right)_{\mu} + \left( \frac{\partial \bar{T}_{(i)}}{\partial t} \right)_{rad} + \left( \frac{\partial \bar{T}_{(i)}}{\partial t} \right)_{WTG} \right] dz \right\} = 0 \quad \text{and} \quad \sum_{i=1,2} \left\{ E_{(i)} - P_{(i)} + L_v \int_{surf}^{z_{top}} \rho \left( \frac{\partial \bar{q}_{(i)}}{\partial t} \right)_{WTG} dz \right\} = 0$$

Sensible heat flux      Heating from condensate      Radiative cooling      Heating from the large-scale circulation      Evaporation minus precipitation      Moistening from the large-scale circulation

**Assumptions:**

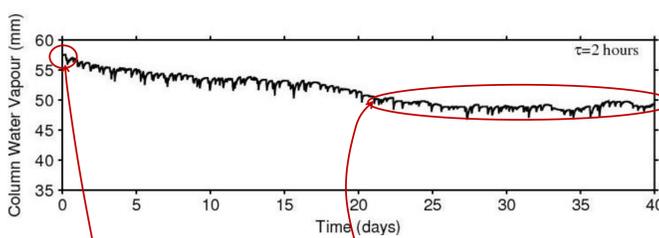
- 1- Horizontal flow with no shear, 2-Large-scale circulation doesn't advect condensate, 3-  $\frac{\partial \bar{\theta}}{\partial z} \gg 1K/km$

## 3. Uncoupled model

- The Met Office Large Eddy Model (LEM) at version 2.4 is used.
- The domain size is  $Y \times Z = 128 \times 20 km$  with horizontal resolution of  $0.5 km$ .
- The lower boundary is a uniform sea surface temperature (SST).
- Radiative cooling is prescribed.
- The model is run for 40 days **with WTG calculations off** and SST of 301.7, 302.7 and 304.7K.
- The last 20 days profiles are averaged to provide the sounding at RCE. Those obtained at 302.7K (**control run**) are used as reference profiles in the reference column model.

## 4. Reference column model

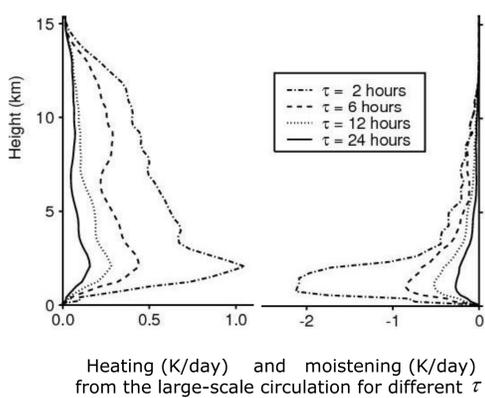
- $\bar{\theta}$  and  $\bar{q}$  are specified in one column (reference column).
- Heat and moisture advected out of the other column (test column) is not received by the reference one: **Heat and moisture budgets are not strictly closed**.
- The test column is initialised with profiles of the reference column.
- The SST in both columns is 302.7K.
- Simulations are performed for  $\tau = 2, 6, 12, 24$  and 120 hours. For gravity waves of mean speed  $50 m/s$ , they correspond to large-scale circulation of horizontal length  $\chi = 360, 1080, 2160, 4320$  and  $21600 km$ .



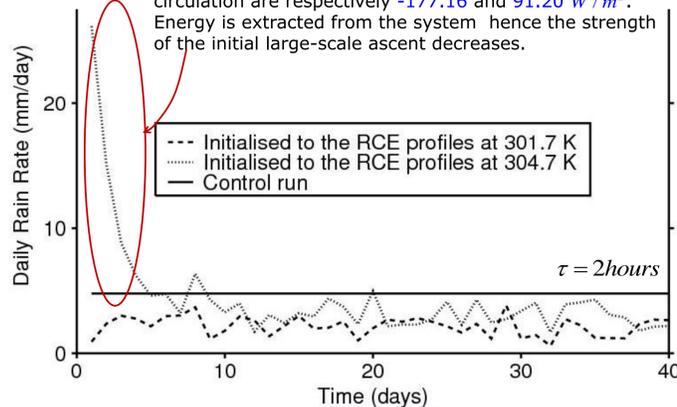
Experiments	Mean rain rate (mm/day)	Evaporation (mm/day)
Control run	4.77	4.80
$\tau = 120$ hours	4.70	4.74
$\tau = 24$ hours	4.40	4.77
$\tau = 12$ hours	4.03	4.78
$\tau = 6$ hours	3.43	4.64
$\tau = 2$ hours	1.99	4.47

Starts from the equilibrium value obtained in the control run      Quasi-steady state **different from the RCE state** obtained in the control run

### Sensitivity to initial conditions

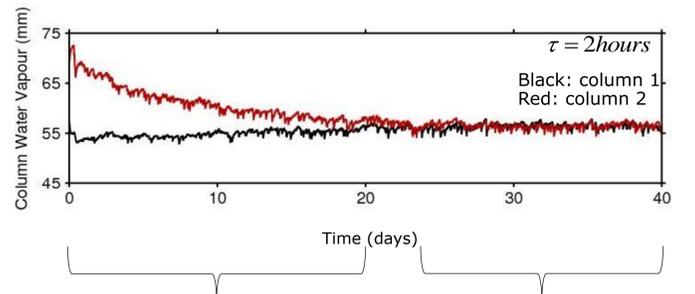


The heating and the moistening from the large-scale circulation are respectively  $-177.16$  and  $91.20 W/m^2$ . Energy is extracted from the system hence the strength of the initial large-scale ascent decreases.



## 5-Coupled Columns Model

- The profiles in neither column are specified.
- Heat and moisture advected out of one column is equal to that received by the other column: **Heat and moisture budgets are closed**.
- Column 1 is initialised to the RCE profiles at 302.7K and column 2 to that at 304.7K.



The model tries to maintain the initialised large-scale circulation but this last for 20 days only.      The two columns adjust to almost the same state which is very similar to the **RCE state** obtained in the control run.

	Mean rain rate (mm/day)	Evaporation (mm/day)
Control run	4.77	4.80
Column 1	4.72	4.73
Column 2	4.85	4.77

The adjustment time scale in the coupled columns model decreases with increasing  $\tau$ .

## 6-Summaries

### Reference column model

- This model has a **unique final state with descent in the test column which does not depend on how it has been initialised**.
- The mean rain rate increases with  $\tau$ .
- The rate of change of evaporation is negligible hence, precipitation variations are mainly controlled by large-scale horizontal moisture advection.

### Coupled columns model

- This new model does not sustain a large-scale circulation no matter the strength of the initial circulation. Hence, **large-scale circulation with descent in the test column is an artefact of the reference column approach**.
- The shorter the value of  $\tau$ , the longer the time required by the model to adjust to an equilibrium with no large-scale circulation.

### Future work

- Examine the equilibrium response of the coupled columns model to inhomogeneous SST.
- Understand how two-way interactions between convection and large-scale circulation influence the transition from shallow to deep convection.

## References

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 Raymond, D., and X. Zeng, 2005: Modelling tropical atmospheric convection in the context of the weak temperature gradient Approximation. *Quart. J. R. Meteorol. Soc.*, 131(608), 1301-1320.  
 Perez, C., A. Sobel, G. Gu, C. Shie, W. Tao, and D. Johnson, 2006: Equilibrium properties of tropical deep convection in cloud-resolving simulations using the weak temperature gradient approximation. *J. Atmos. Sci.*

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